Oceanic and Atmospheric Drivers of



Post-El-Niño Chlorophyll Rebound in the Equatorial Pacific



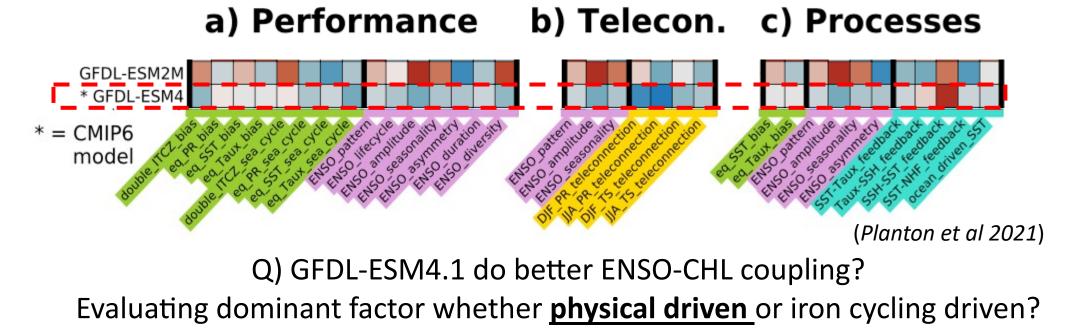
SCRIPPS INSTITUTION OF

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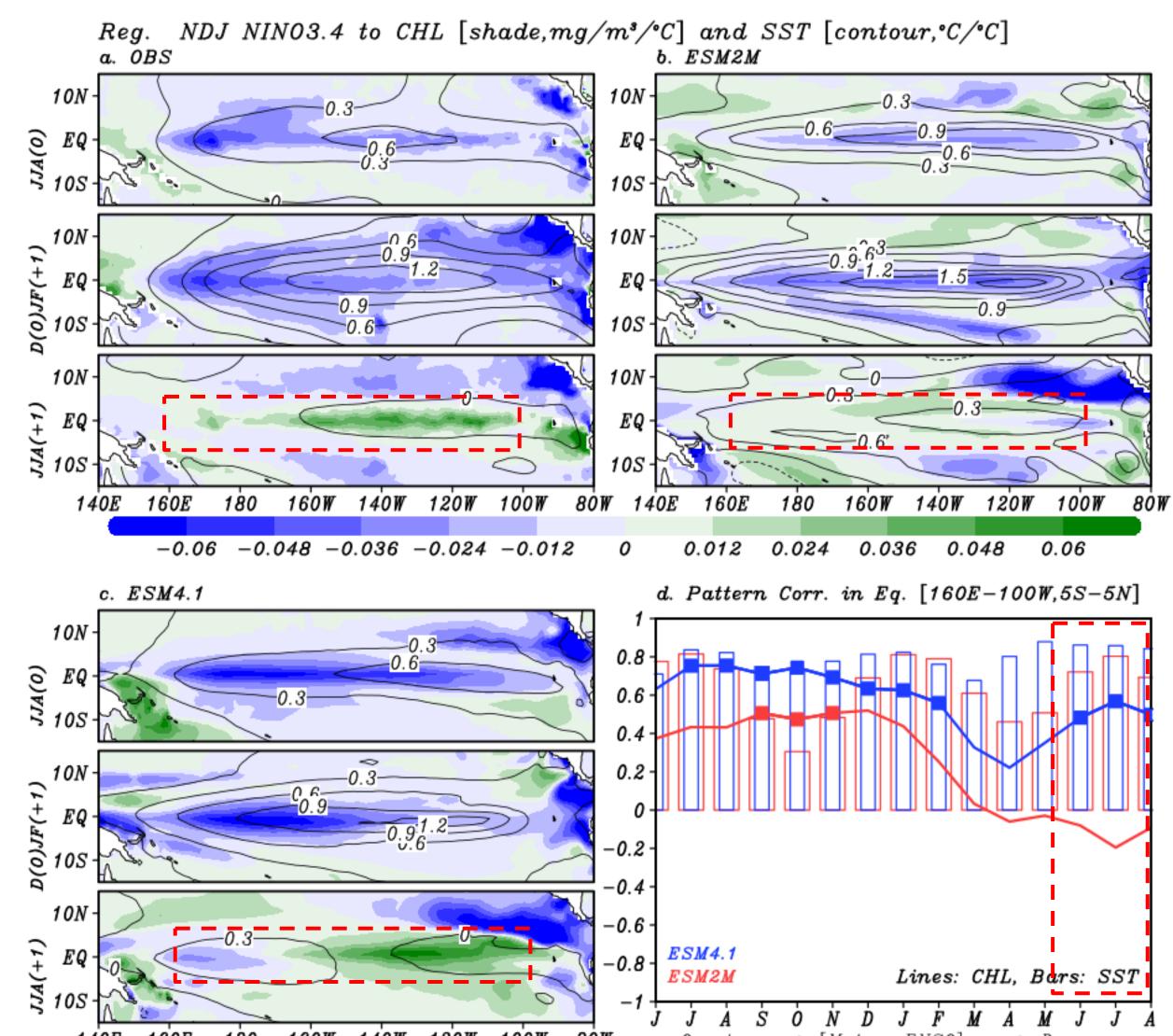
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Motivation

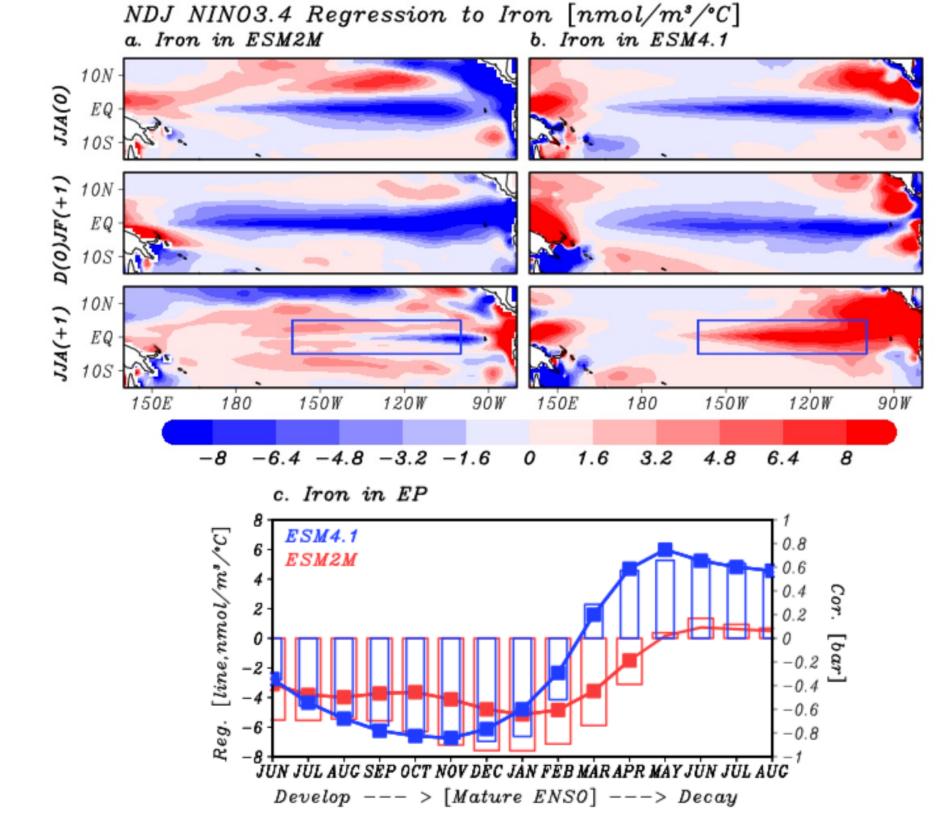
- ✓ New generation GFDL-ESM4.1 (*Dunne et al 2020; Stock et al 2020*)
 - Better fidelity of ENSO simulation (*Planton et al 2021*)
 - > Dynamic dust-iron deposition in land to ocean delivery (*Evans et al 2016*)



ENSO-CHL coupling in OBS, ESM2M, ESM4



Post El Niño Iron rebound



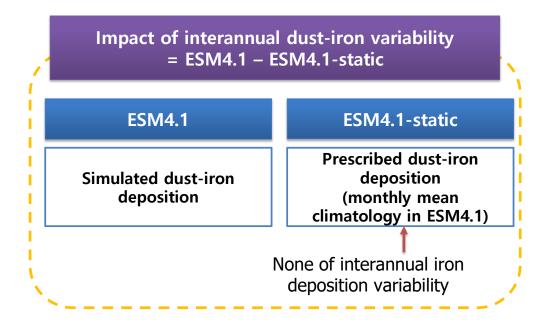
Datasets

- piControl simulations from two generations of GFDL earth system models: years 1001–1100 of ESM2M; years 501–645 of ESM4.1
- Satellite-derived ocean color for chlorophyll estimation: the European Space Agency Ocean Colour Climate Change Initiative project (ESA-CCI) version 4.2
- NOAA Extended Reconstructed SST version 5 (ERSST v5).
- Vertical temperature profiles: version 3.4.2 of the Simple Ocean Data Assimilation (SODA)
- Two satellite-derived aerosol optical thicknesses: NOAA Climate Data Record (CDR) of AVHRR Daily and Monthly Aerosol Optical Thickness (AOT) over Global Oceans, Version 3.0; SeaWiFS Deep Blue Level 3 monthly product which contains monthly global gridded (1° × 1°) data derived from SeaWiFS Deep Blue Level 3 daily gridded data (Hsu et al., 2013).

Methods

- ✓ Linear regression for variables against November to January (NDJ) averaged SST anomalies (SSTa) in the Niño3.4 region (170° −120° W, 5° N−5° S; Niño3.4 SSTa) (Trenberth, 1997) is employed to diagnose general ENSO influences.
- \checkmark Converted to 1degree horizontal resolution
- ✓ Observation for September 1997 until December 2018
- Linear trends and seasonal cycles were removed from all data sets.
- ✓ All data sets are smoothed using a 3-month running mean to focus on the seasonal ENSOchlorophyll evolution and its variability.

ESM4.1-static run



*Integration - ESM4.1: 145-yr (501-645yrs) of CMIP6 picontrol - ESM4.1-static : 145-yr runs (501-645yrs) from 499yr initial condition of ESM4.1

> 0.8 0.64 0.48 0.32 0.16

> > -0.16 -0.32 -0.48

-0.64 -0.8

****Oceanic Driver**

Reg. NDJ NINO3.4 at vertical section of Eq. Pacific (2°N-2°S) a. Temp. in ESM2M b. Temp. in ESM4.1 c. Temp. in SODA3.4.2 [°C/°C] 140E 160E 180 160W 140W 120W 100W 80W Onset --- > [Mature ENSO] ---> Decay

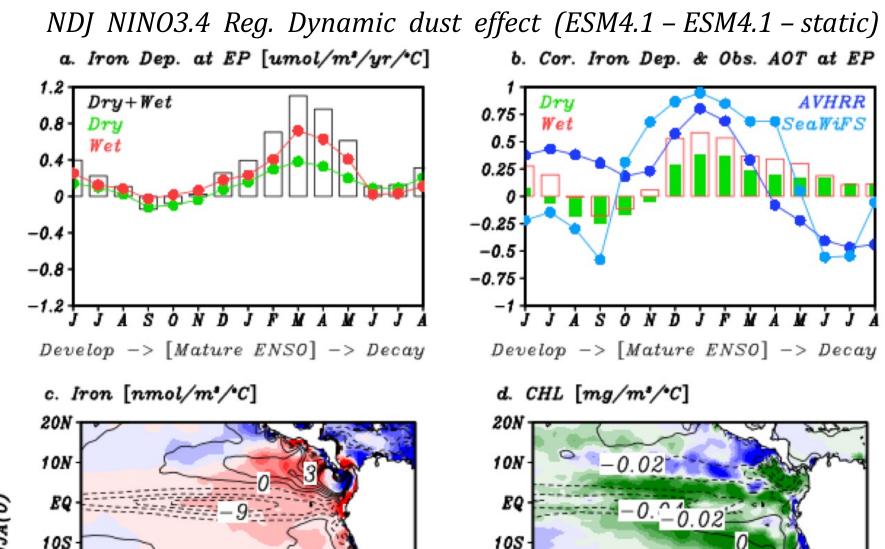
Figure 1. Spatially resolved ENSO-regression of temperature (contours, ° C/° C) and chlorophyll concentration (shading, mg/m3/° C) anomalies against ENSO in (a) satellite, SODA, (b) ESM2M, and (c) ESM4.1. Values are shown for ENSO onset (JJA[0]), mature state (D[0]JF[+1]), and decay (JJA+1) ENSO states; (d) pattern correlations of ESM2M (red) and ESM4.1 (blue) regression coefficients (i.e., panels b and c) with SODA (bars) and satellite chlorophyll (lines; square symbols denote the Student's t-test statistical significance at 95% confidence level) regression coefficients (i.e., panel a) in the EP (160° E–100° W, 5° N–5° S).

Mechanism: El Niño to dust-iron

Westerly wind burst and El Niño onset

Figure 2. ENSO-regressed patterns in spatially resolved surface iron concentration anomalies (nmol/m3/° C) in (a) ESM2M and (b) ESM4.1. (c) lead-lag ENSO regression (lines; square symbols denote the Student's t-test statistical significance at 95% confidence level) and correlation (bar) coefficients in regional monthly iron concentration anomalies in the eastern EP (160° -100° W, 5° S–5° N).

Atmospheric Driver



anomalies in ESM4.1-static. ENSO-regressed (e) iron and (f) chlorophyll anomalies

averaged in the eastern EP in ESM4.1 (blue) and ESM4.1-static (sky blue).

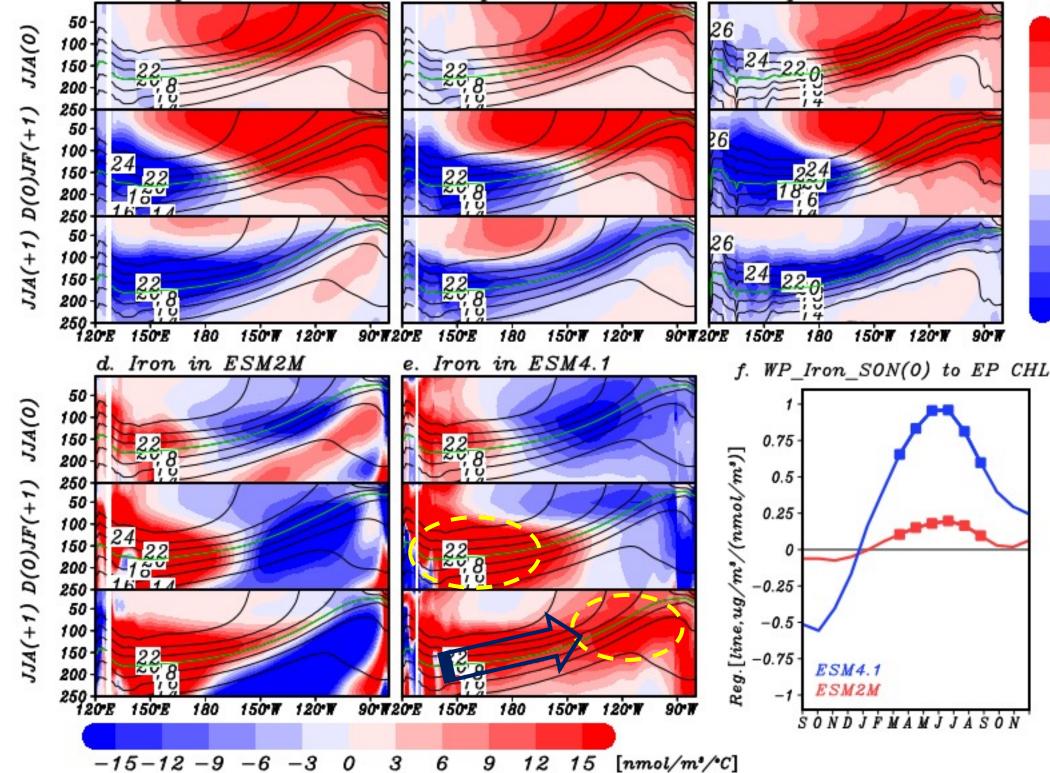
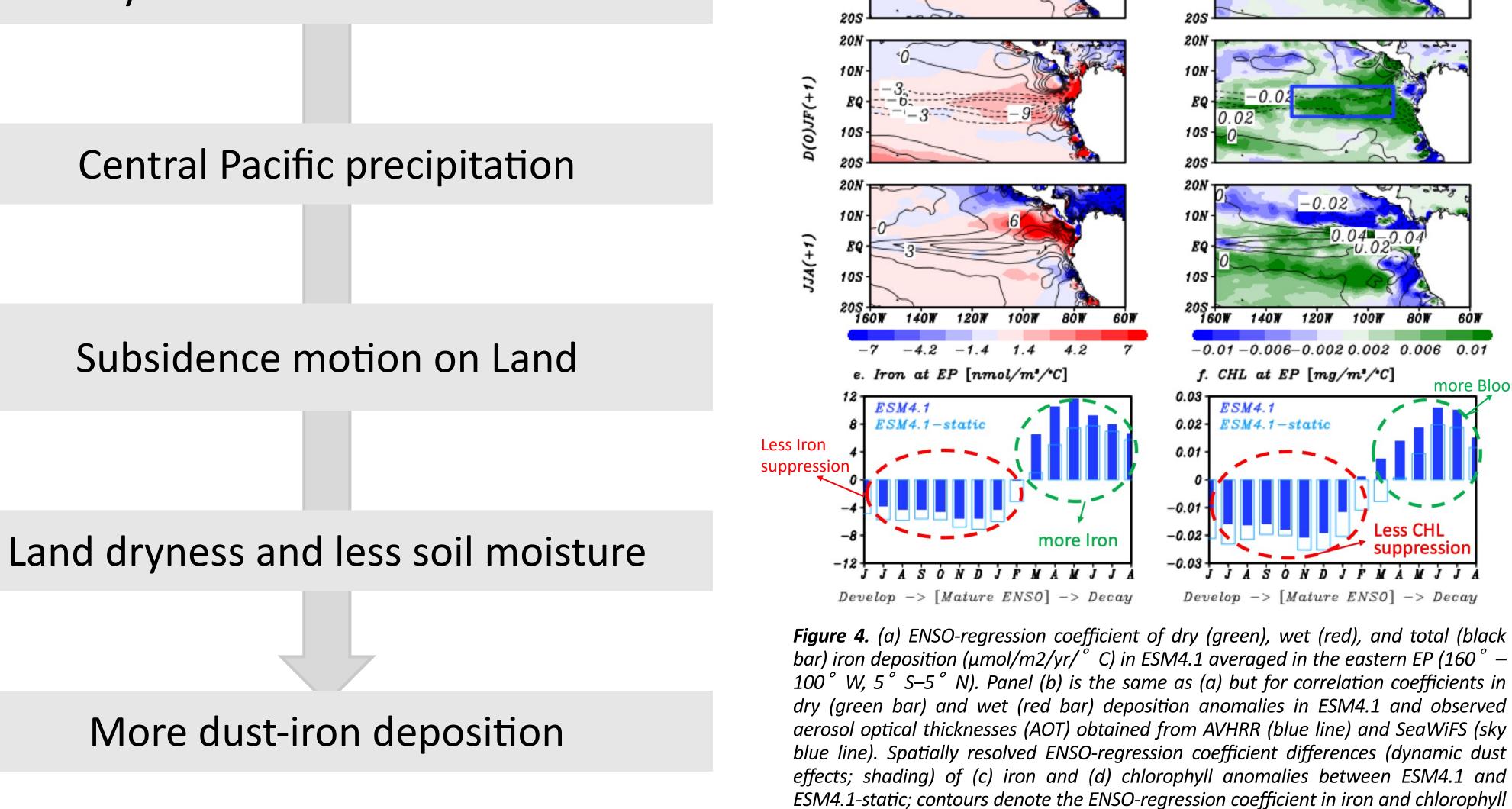


Figure 3. ENSO-regressed coefficients of vertical sections of equatorial temperature anomalies in (a) ESM2M, (b) ESM4.1, (c) SODA and dissolved iron concentration anomalies (shading) in (d) ESM2M, (e) ESM4.1. The temperature climatology from each model and SODA has been overlaid (contours) to aid interpretation and 20° C isotherm (green line) at equator $(2^{\circ} N-2^{\circ} S)$. (f) Regressed coefficients (square symbols denote the statistical significance at 95% confidence level) of iron concentration anomalies in the WP subsurface region averaged in $(130^{\circ} -160^{\circ} E, 2^{\circ} N-2^{\circ} S, 100-200 m depth)$ at SON(0) against the eastern EP surface $(160^{\circ} -100^{\circ} W, 5^{\circ} S-5^{\circ} N)$ monthly chlorophyll concentration anomalies simulated in ESM2M (red) and ESM4.1 (blue).



Conclusion

Better capturing post-El Niño CHL rebound pattern in GFDL-ESM4.1 (Fig. 1)

- Much more iron supply (Fig. 2) in post-El Niño event into the eastern equatorial Pacific in ESM4.1 than ESM2M
- **Oceanic driver (Fig. 3): Iron trapping and its propagation via equatorial undercurrent (Dominant factor to ENSO-CHL rebound)
- Atmospheric driver (Fig. 4): Less precipitation on land increases dust-iron deposition (Argument factor to ENSO-CHL sensitivity)
 - > The chlorophyll rebound in GFDL-ESM4.1 may provide a key source of resilience to marine ecosystem simulation in the

equatorial Pacific for management of living marine resources.

Further descriptions and references at:

Lim, H.-G., Dunne, J. P., Stock, C. A., Ginoux, P., John, J. G., & Krasting, J. (2022). *Oceanic and atmospheric drivers of post-El-Niño chlorophyll rebound in the equatorial Pacific. Geophysical Research Letters*, 49, e2021GL096113. <u>https://doi.org/10.1029/2021GL096113</u>