

Multivariate analysis of wind stress and curl over the Japan/East Sea, based on satellite scatterometry data

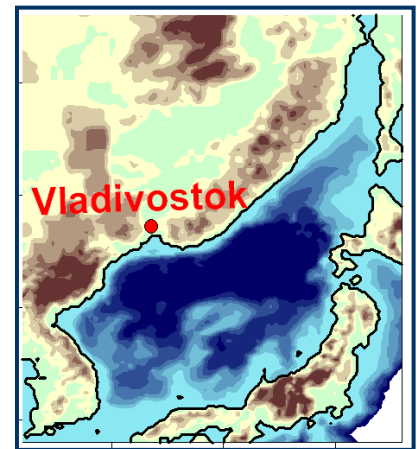
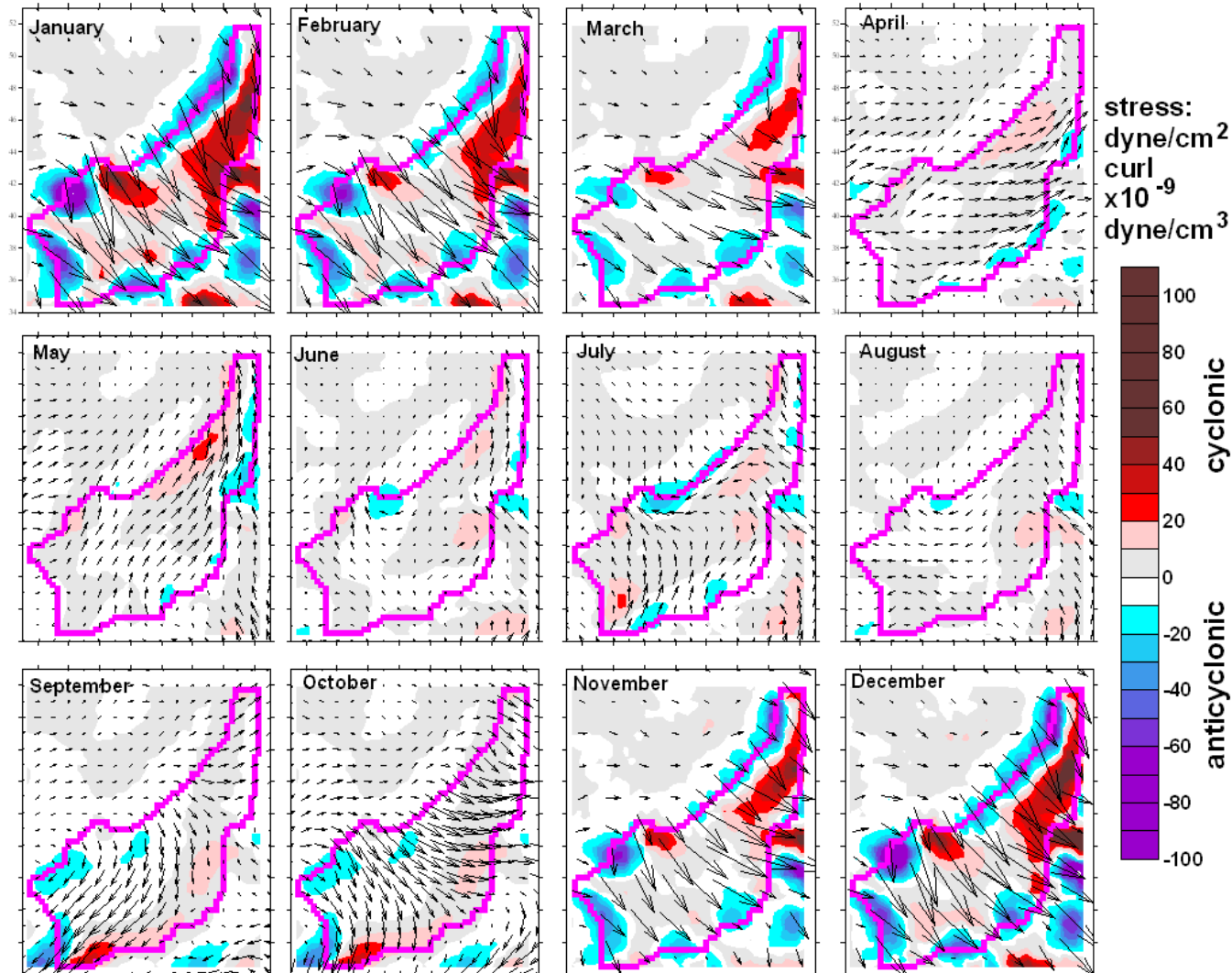


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PICES 2010, October 22–31, 2010, Portland, Oregon, USA

Motivation

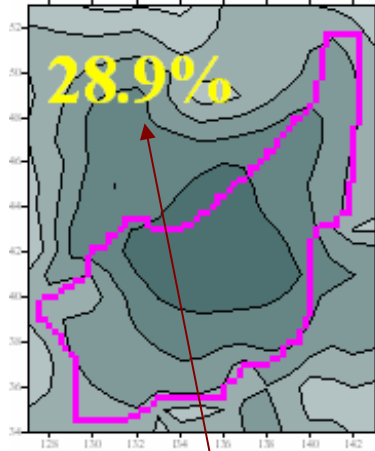
Monthly mean wind stress and curl in the JES area



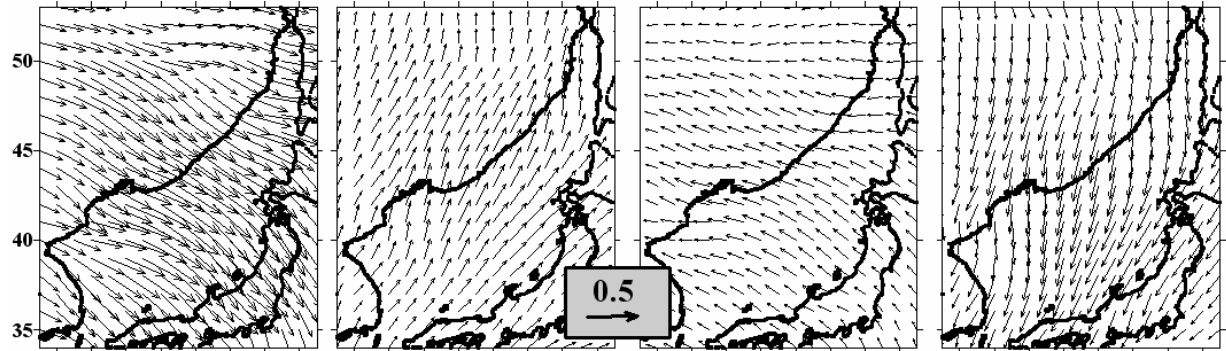
Stable conditions
in winter,
changeable and
weak, on
monthly time
scale, wind in
summer →
Are there any
patterns in
summer wind?

Previous results from 1°-gridded NCEP reanalysis (ancillary data for the SeaWiFS Project)

Mode 1

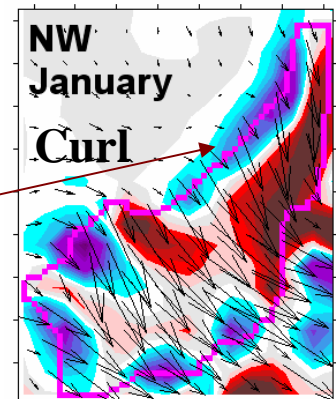
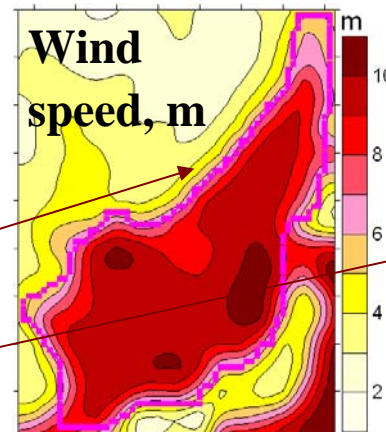


Leading complex EOF mode from wind vectors accounted for the general wind direction → effect of the East Asia monsoon (Trusenkova et al., 2009).



However,
fraction of the total variance was
not high → any impact of higher
modes?

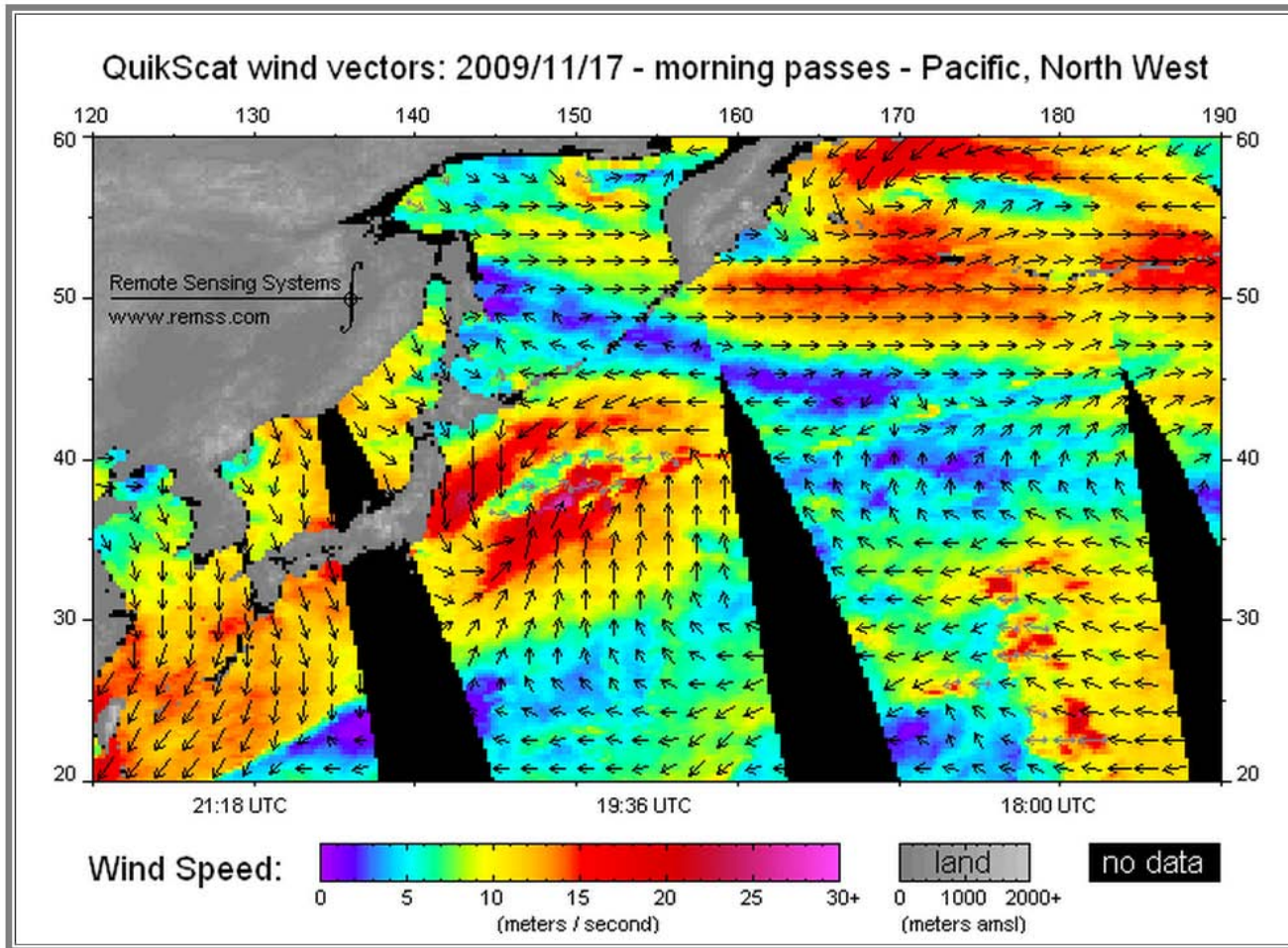
There were fictional alongshore
wind gradient zones →
fictional curl zones.



Purpose

To revisit the spatial/temporal patterns of wind stress and curl over the JES by multivariate statistical analysis of satellite scatterometry data

Sea winds from QuikSCAT



High resolution data over the water surface.

However, frequent gaps due to satellite swath divergence and rain contamination →

QSCAT merged with reanalysis provides data over the sea and land.

Data

QSCAT/NCEP Blended Ocean Winds

from Colorado Research Associates,

July 1999 - July 2009,

34°-53°N, 127°-143°E (JES and adjacent land),

6 hours, 0.5° grid,

1287 wind stress boxes and

386 stress curl boxes (over the sea only),

14724 times.

Techniques of EOF analysis

1) The complex form of the EOF analysis applied to wind vectors.

$X(r, t) = \sum A_k^*(r)B_k(t)$, where

$X(r, t) = U(r, t) + iV(r, t)$, U/V are zonal/meridional wind components,

$A_k(r) = A_k(r)e^{-i\phi}$ are spatial CEOFs and $*$ denotes complex conjugate,

$B_k(t) = B_k(t)e^{-i\phi}$ are principal components (PCs),

A_k, B_k are spatial and temporal amplitudes \sim mode intensity,

ϕ_k, ϕ_k are spatial and temporal phases ($-180^\circ, 180^\circ$) \sim wind shear in space and time.

2) Correlations \rightarrow detection of low amplitude signals \sim wind over the land

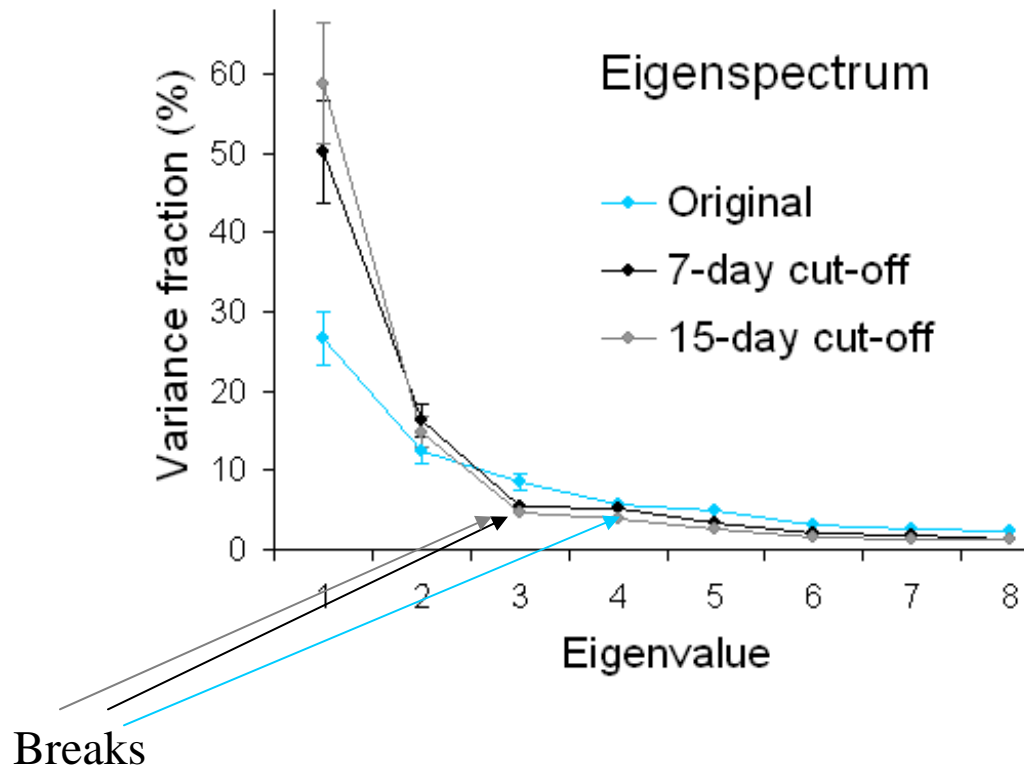
Covariances \sim signal magnitude \rightarrow anomalies from weaker signals can be lost.

3) Low-pass filtering \rightarrow by the inverse wavelet transform,

Morlet mother wavelet of the 6th order.

Decompositions of wind stress vectors

- Original dataset.
- Low-pass filtered dataset, with the 7- and 15-day cut-off periods.



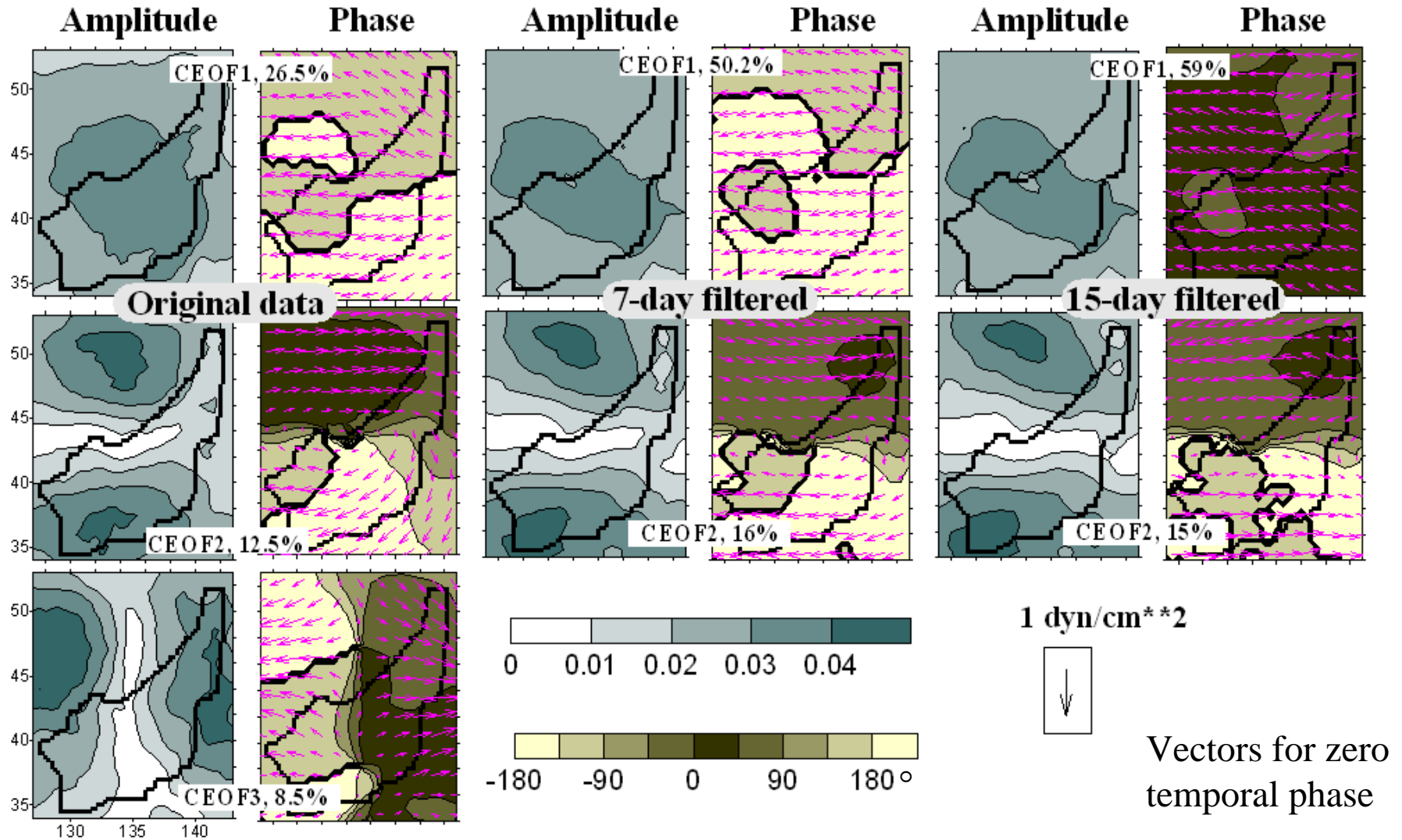
Modes 1 – 3
from the original data
and
modes 1 & 2
from the filtered data
pass the Montecarlo test
and are non-degenerate in the
sense of errors.

High fraction of variance
(>50%) for mode 1 from the
filtered data.

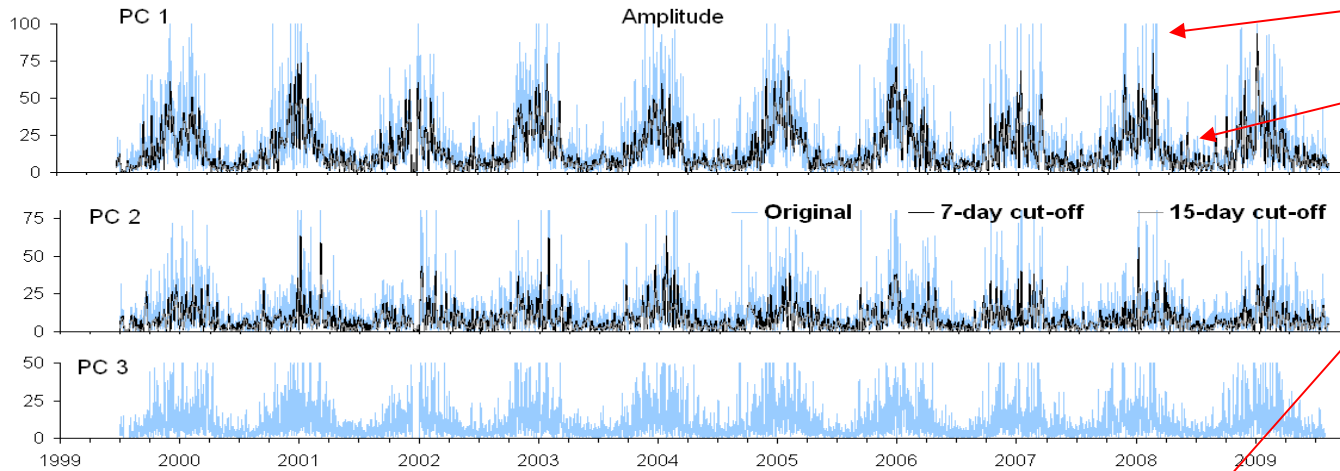
Similar variance fraction for
mode 2.

CEOF modes of wind stress vectors: spatial patterns

Similar patterns of modes 1 & 2 from the original and filtered data.

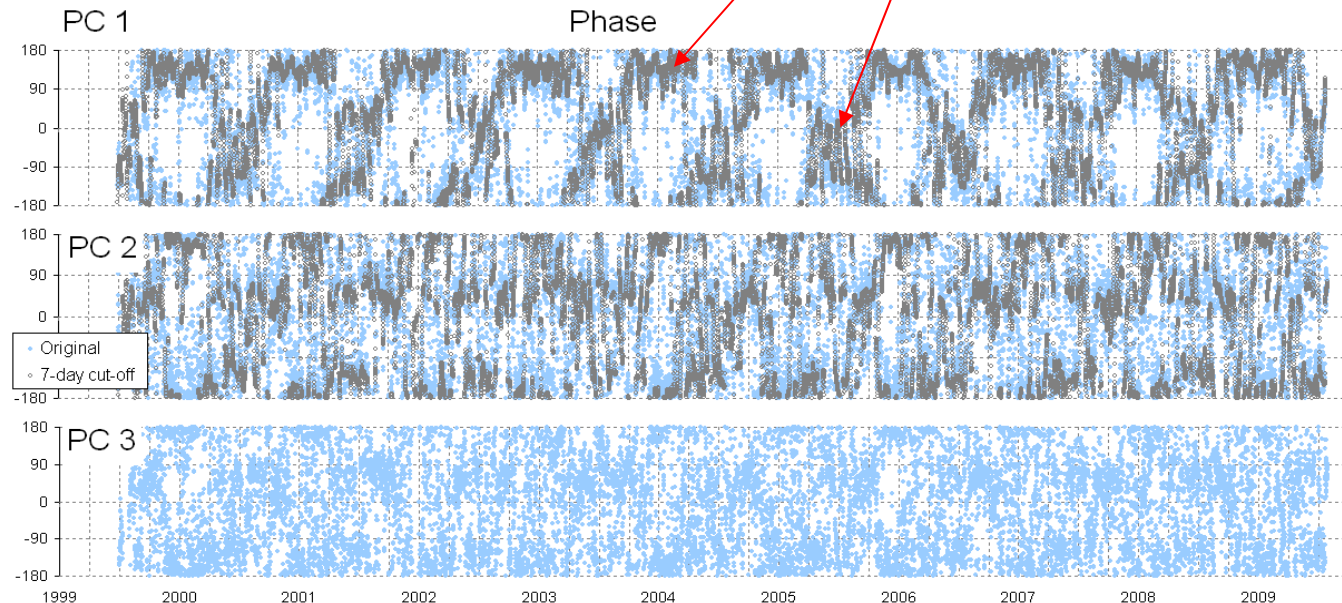
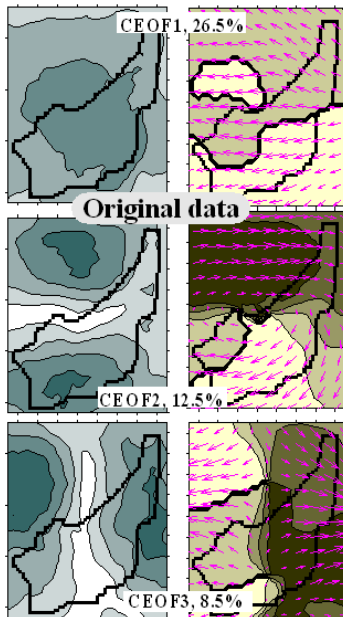


CEOF modes of wind stress vectors: temporal patterns



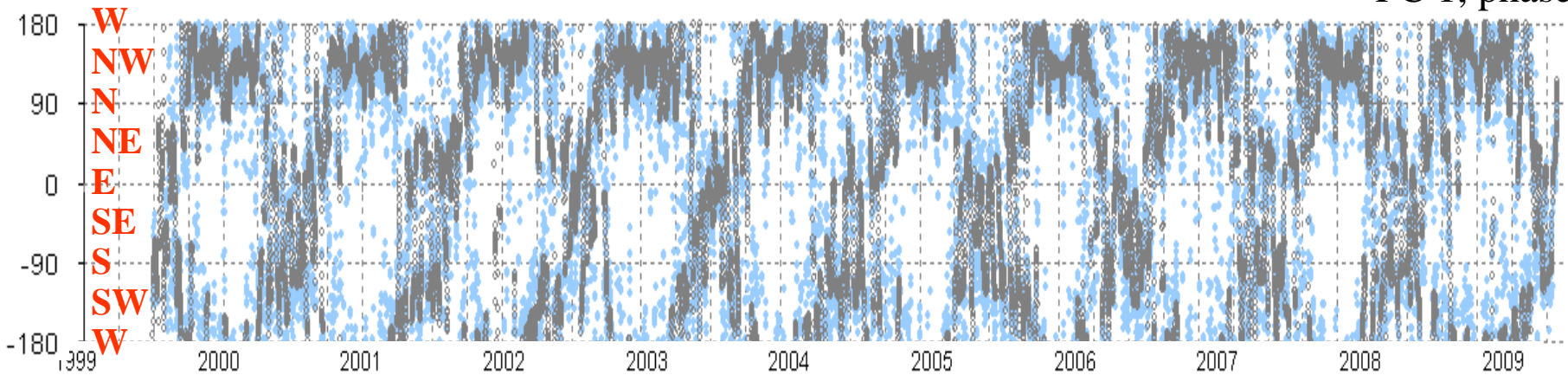
Winter:
strong wind
Summer:
weak wind

Winter:
Persistent NW wind
Summer:
changeable wind



Mode 1: general wind direction

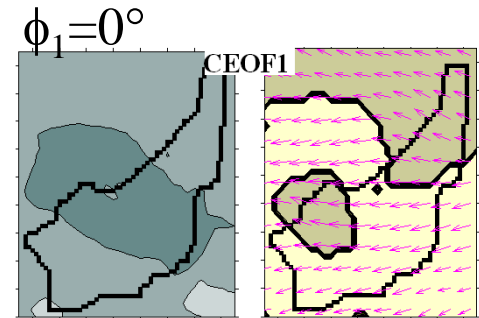
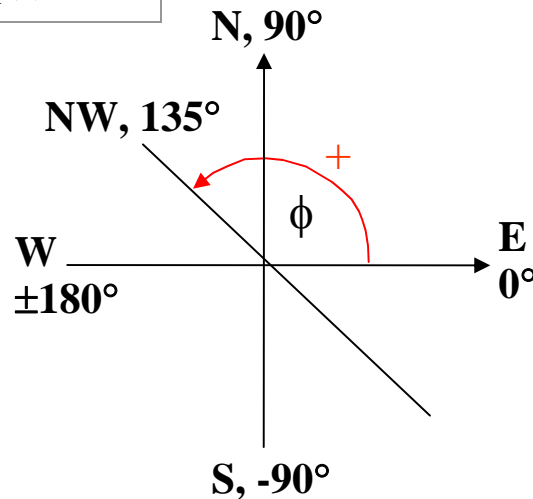
PC 1, phase



$$X_k(r,t) = (A_k(r)e^{-i\phi(r)})^* B_k(t)e^{-i\phi(t)}$$

$\phi_t = 0^\circ \sim$ eastern wind

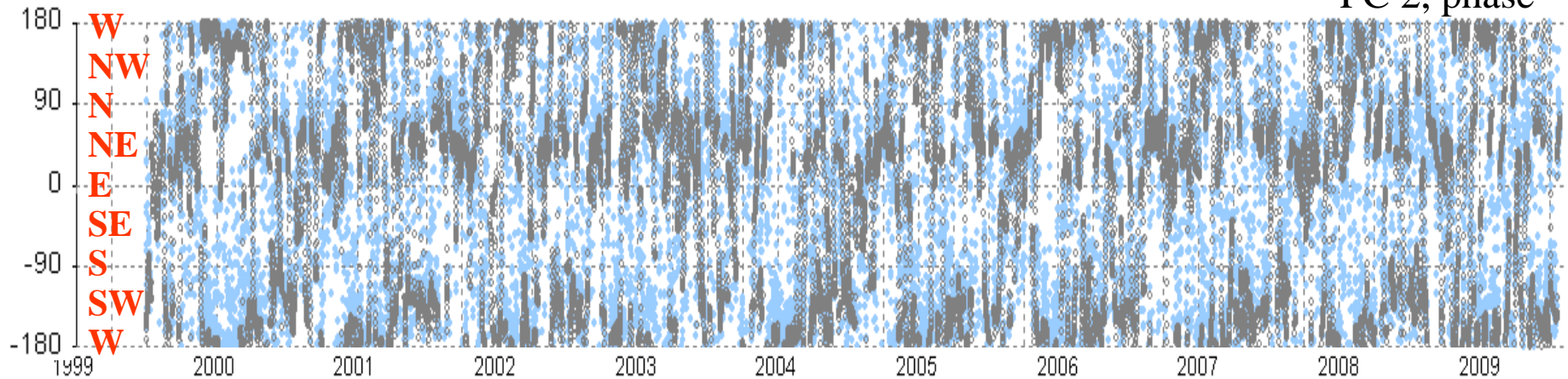
$\phi_1 = 135^\circ \sim$ NW wind



Noisy original $\phi(t)$,
patterns of seasonal
wind shifts in the
filtered data.

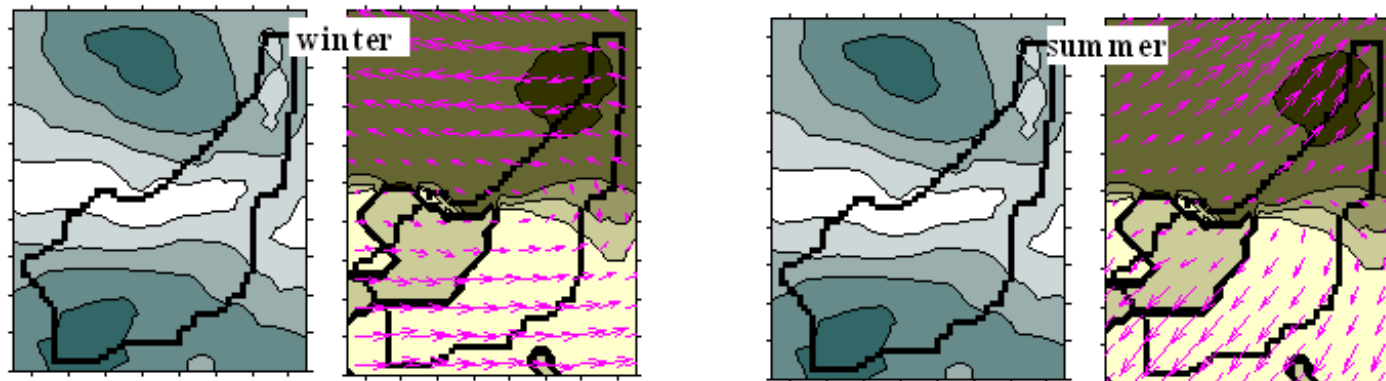
Mode 2: wind seesaw over the southern and northern JES

PC 2, phase



$\phi = 0 \sim$ eastern wind

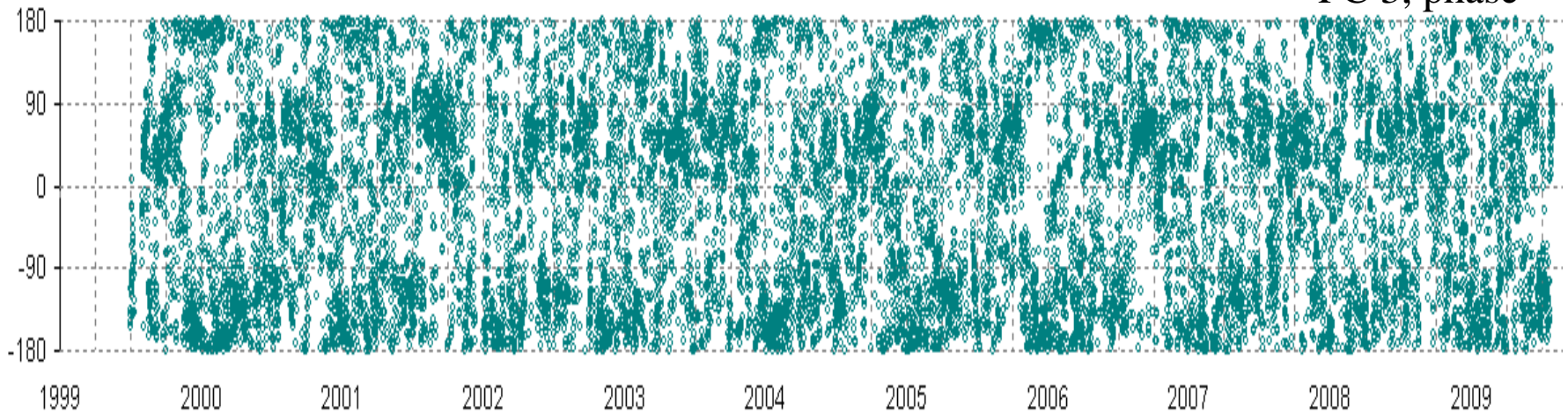
Typical contribution of mode 2 from the filtered data set



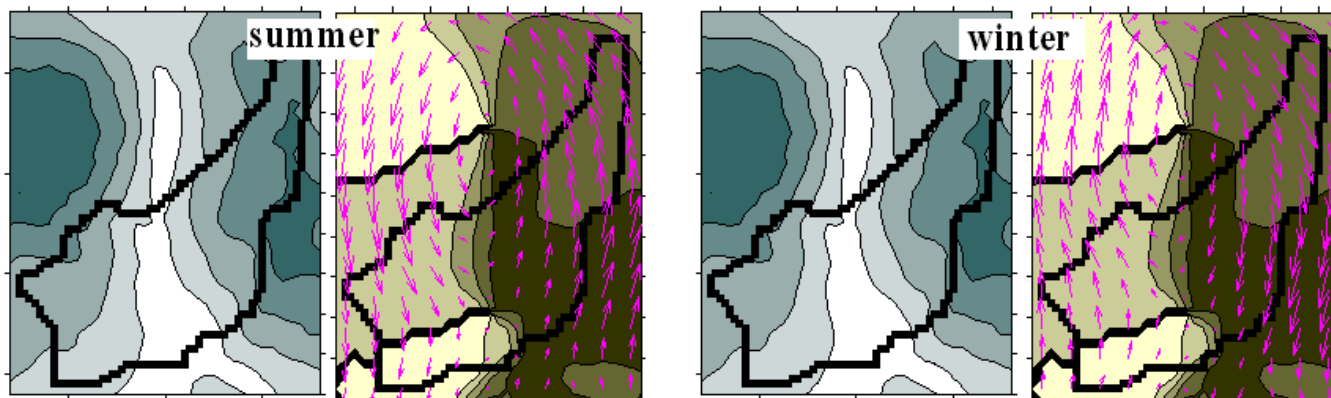
Mode 3: cyclonic / anticyclonic vortex

(original fields only)

PC 3, phase



Typical contribution of mode 3

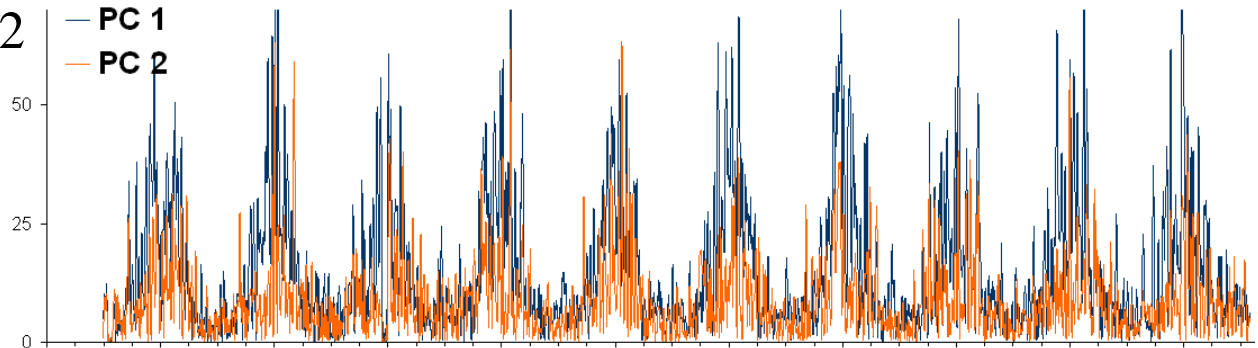
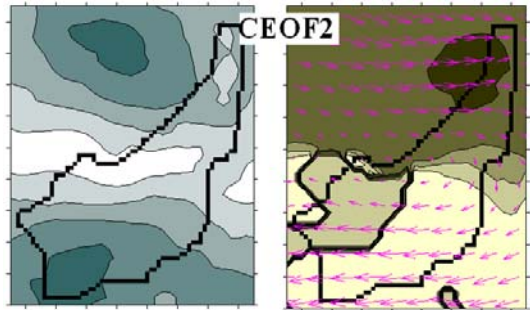
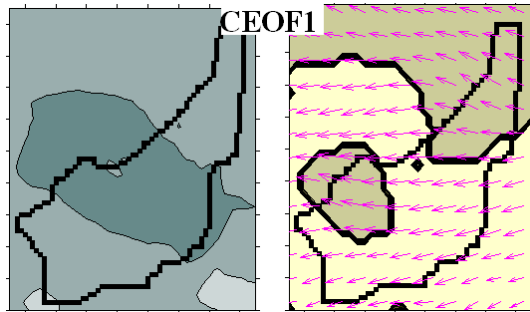


Cyclone

Anticyclone

General wind direction: any contribution from the higher modes?

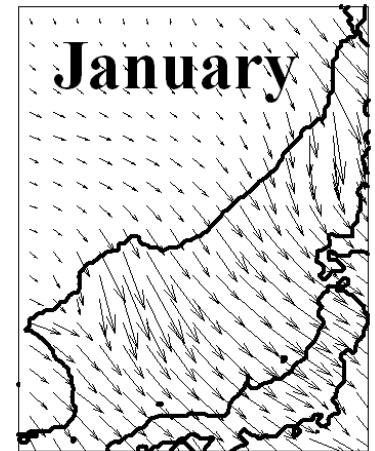
Wind shift due to CEOF 2



Temporal amplitude: $PC1/PC2=1.5$.

Wind shift:

$< 10^\circ$ in the northern JES
and $< 20^\circ$ off the Tsushima Strait



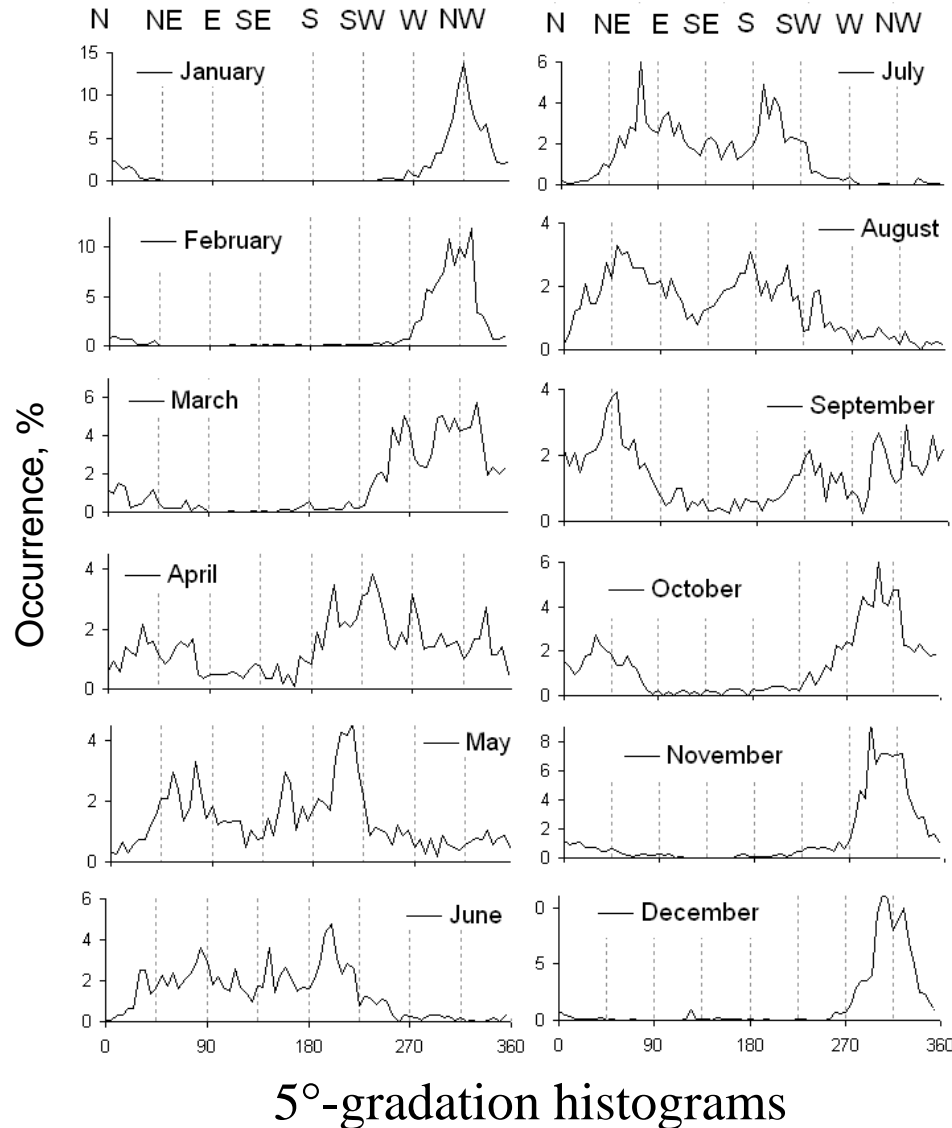
Mode 3 does not affect the wind shifts beyond the synoptic scale

General wind directions over the JES can be deduced from the leading mode alone –

The East Asia Monsoon Mode

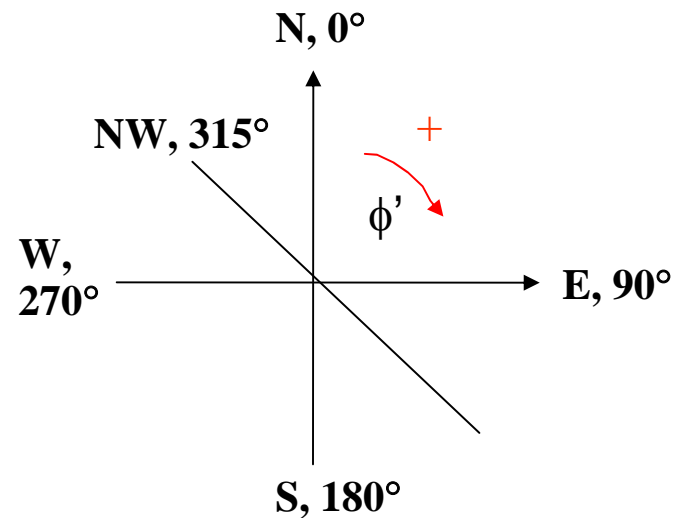
Monthly occurrence of the monsoon winds

(from the temporal phase of mode 1 from low-pass filtered data)



Phase transform

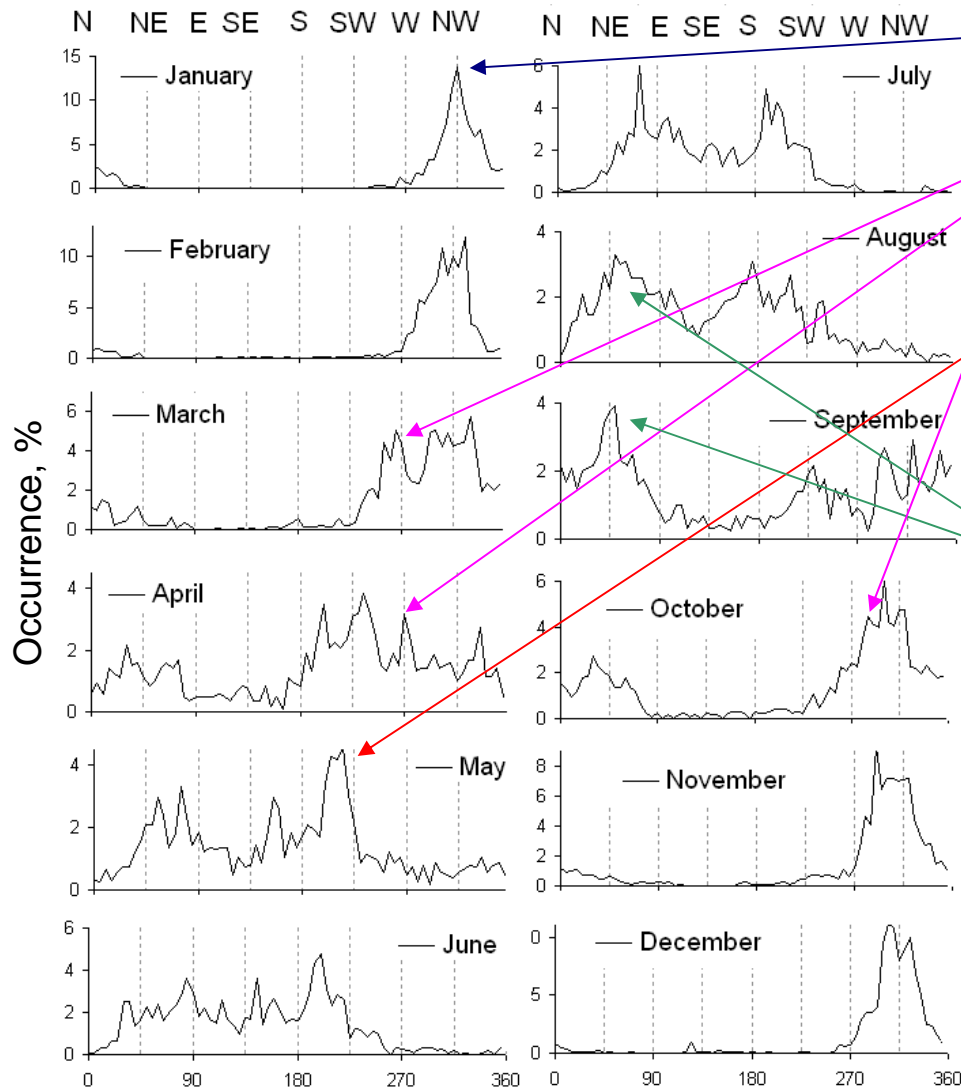
$$\phi' = -\phi + 90^\circ$$



$\phi' = 0^\circ \sim$ northern wind

$\phi' = 315^\circ \sim$ NW wind

Monthly occurrence of the monsoon winds



NW: dominant in November through February;

W: separate mode in March and April and merged with NW in October;

SW: occurs in April through September, strongest in May and June;

Easterly modes: occur more frequently in late summer;

NE: strong in August and September.

November – February: the only mode;

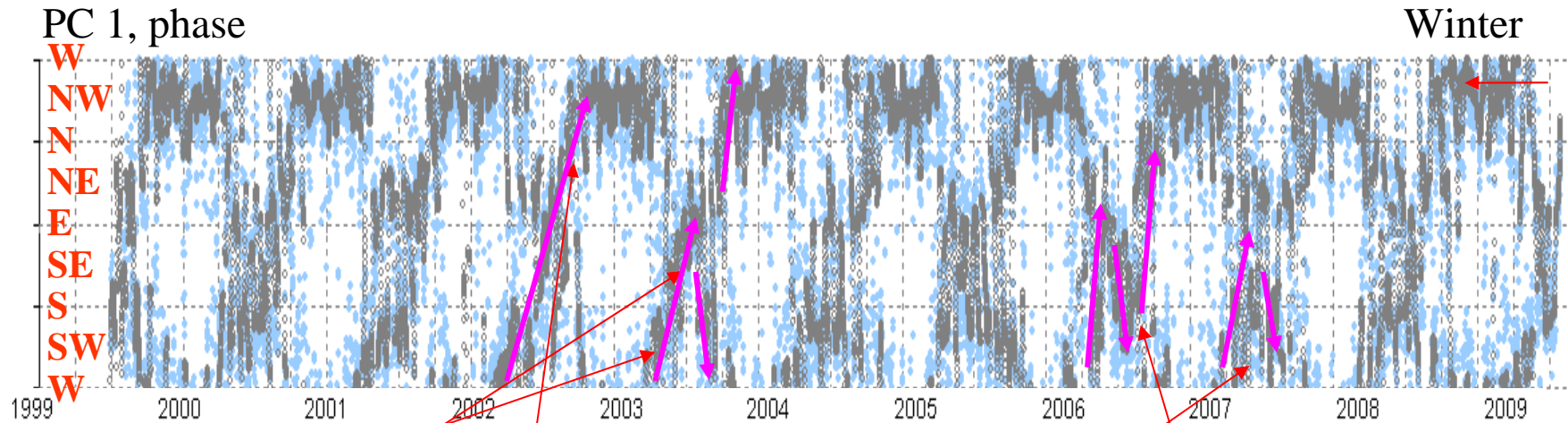
April and September: fuzzy;

July: bimodal (SW and E).

The southern wind is never a dominant mode.

5°-gradation histograms

Seasonal wind shifts



Prevailing transitions:

Spring: NW → W → SW

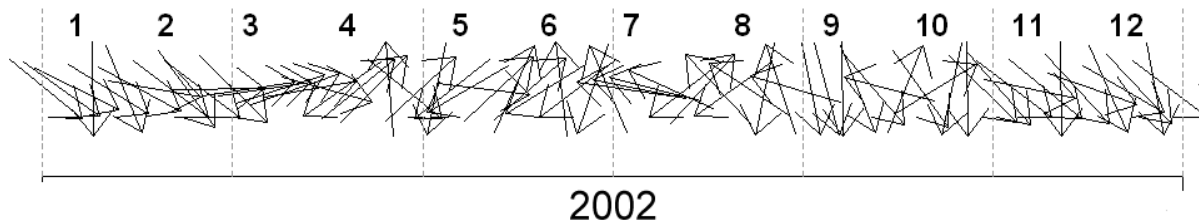
Summer: SW → SE → E

Autumn: E → NE → N → NW or E → SE → SW → W → NW

Strong intraseasonal shifts beyond the synoptic time scale

Time series of monsoon winds (beyond the synoptic time scale)

$$X_1(t) = (A_1^{\text{av}} e^{-i\phi_1^{\text{av}}}) * B_1 e^{-i\phi_1(t)}$$



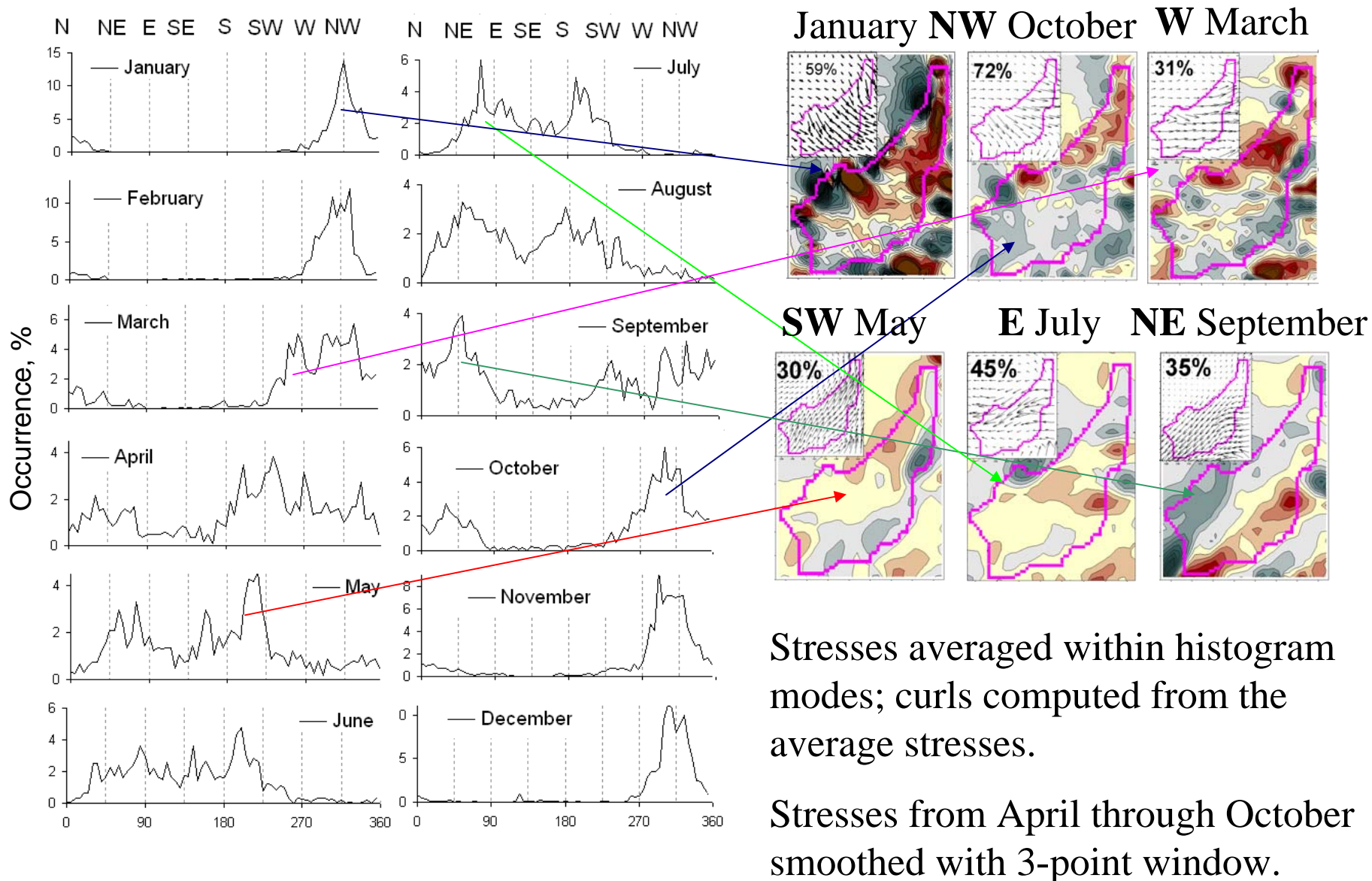
Autumn: E → NE → N → NW



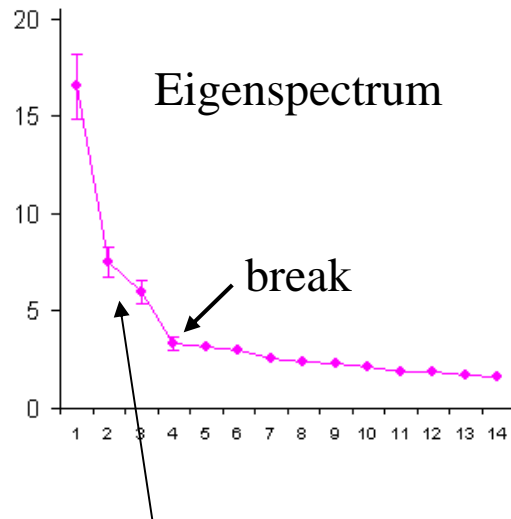
Autumn: E → SE → SW → W → NW

Intraseasonal variability beyond the synoptic scale (1 vector for 5 days shown)

Typical patterns of wind stress and curl



Decomposition of wind stress curl



EV 2 and 3 are close,
followed by a break

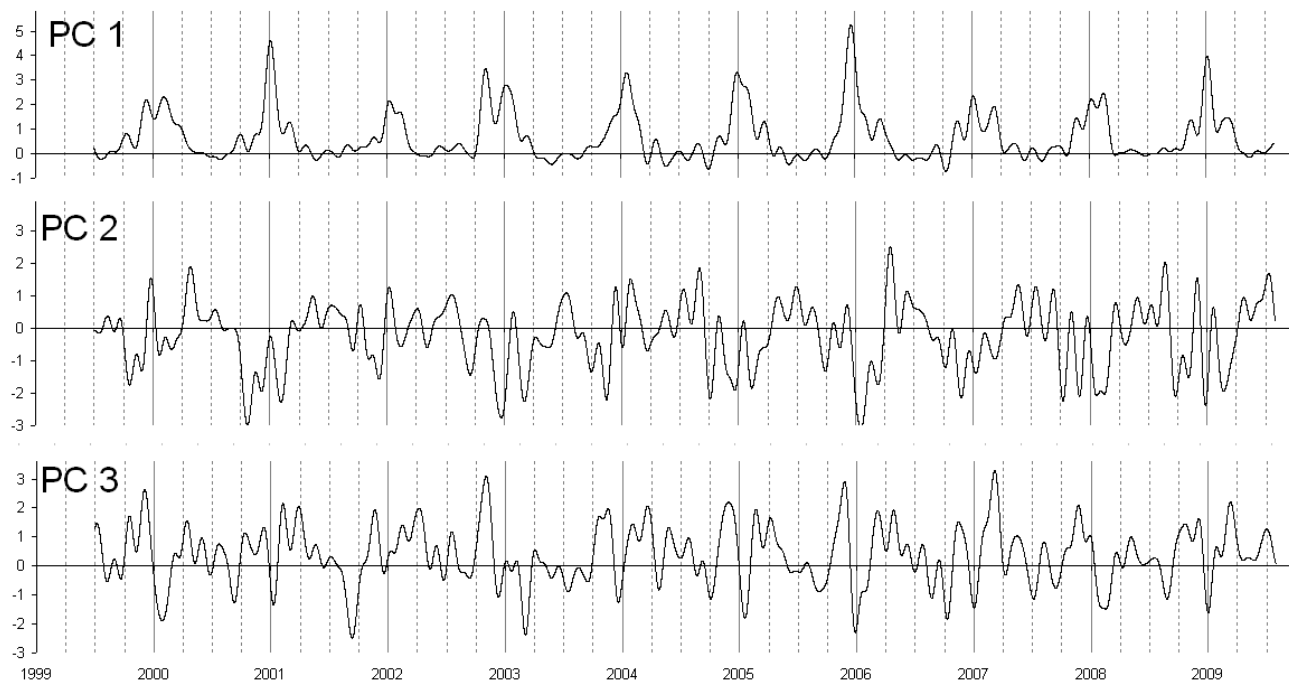
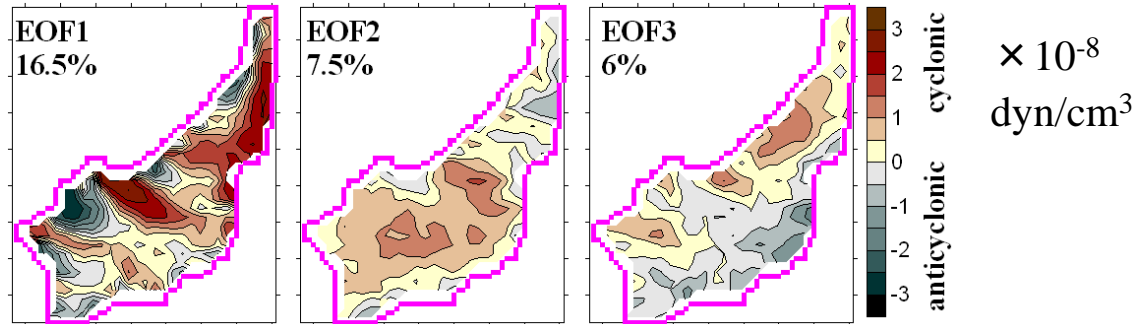
Low-pass filtered wind stress curl fields,
with the 40-day cut-off period.

Modes 1-3 are statistically significant

Mode 1 can be considered separately

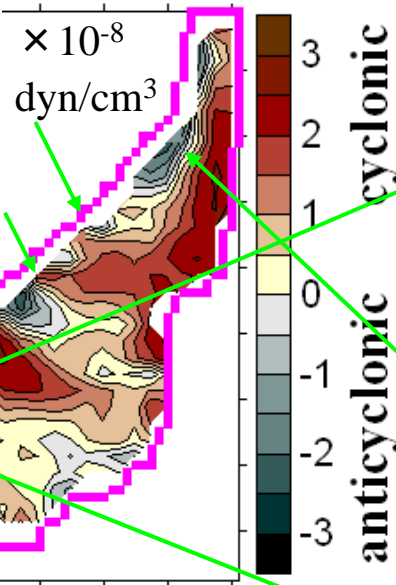
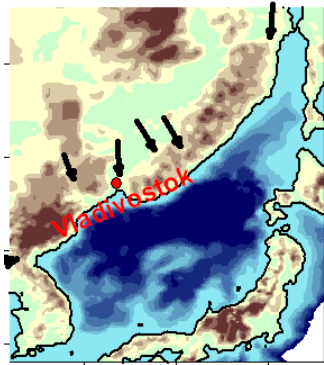
Modes 2 and 3 are better considered together

EOF modes of low-pass filtered wind stress curl

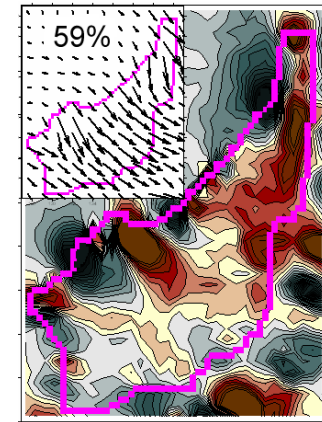


Winter mode

Topographic gaps



Cyclonic curl;
curl dipole
off Vladivostok
(Kawamura and
Wu, 1998),
several dipoles near
the western coast,
AC curl zone off
Khabarovsk Krai
coast,
cyclonic curl zone
off the
East Korea Bay.



January, stress
from histogram



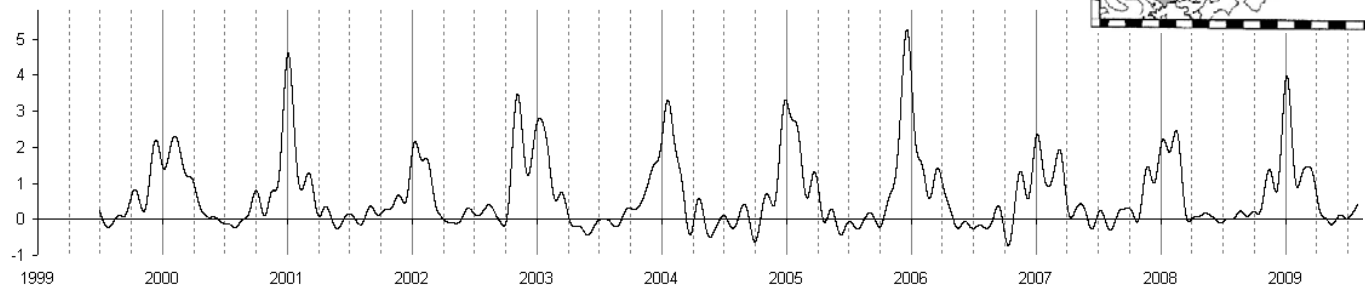
The EKWC separation from the coast

Strong mode:

winters 2000/2001,
2005/2006

Weak mode:

winters 2001/2002
2006/2007,
2007/2008



PC1 is above zero in November through February (mid March)

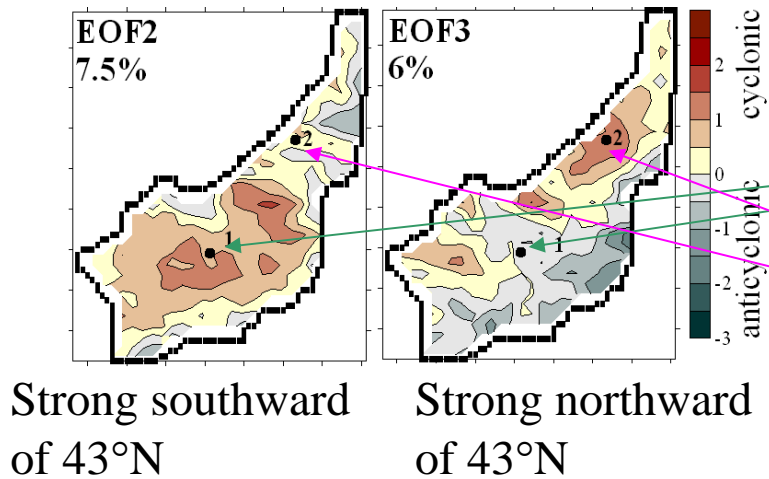
Warm season mode: combined effect of EOF 2 and EOF 3 (significant when PC 1 is close to zero)

Contribution of modes 2 and 3

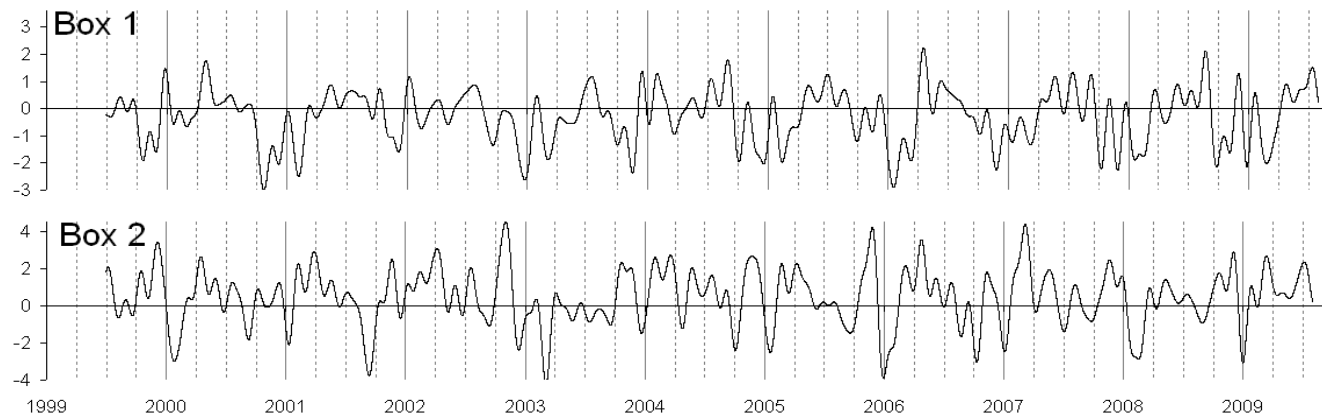
$$X_{1,2}(t) = A_2(r_{1,2})B_2(t) + A_3(r_{1,2})B_3(t)$$

Box 1 (central JES): EOF 2 dominates,
 $\text{corr}(\text{PC2}, X_1) = 0.98$

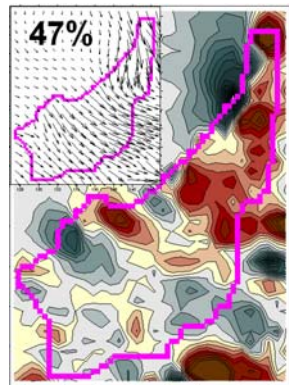
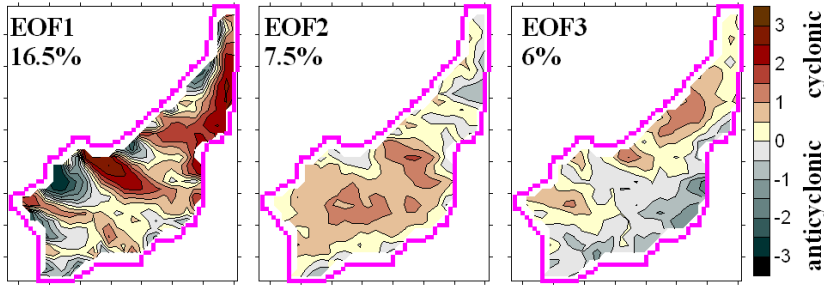
Box 2 (northeastern JES): EOF 3
dominates, $\text{corr}(\text{PC3}, X_2) = 0.97$



Curl ($\times 10^{-8}$ dyn/cm³)

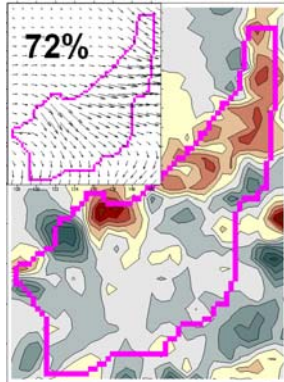


Warm season mode: the central JES

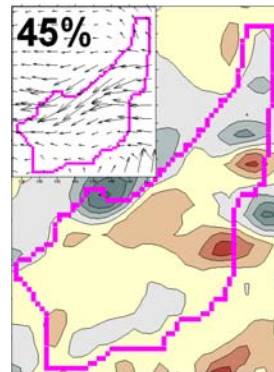


March

NW wind in March
and October: AC
curl



October

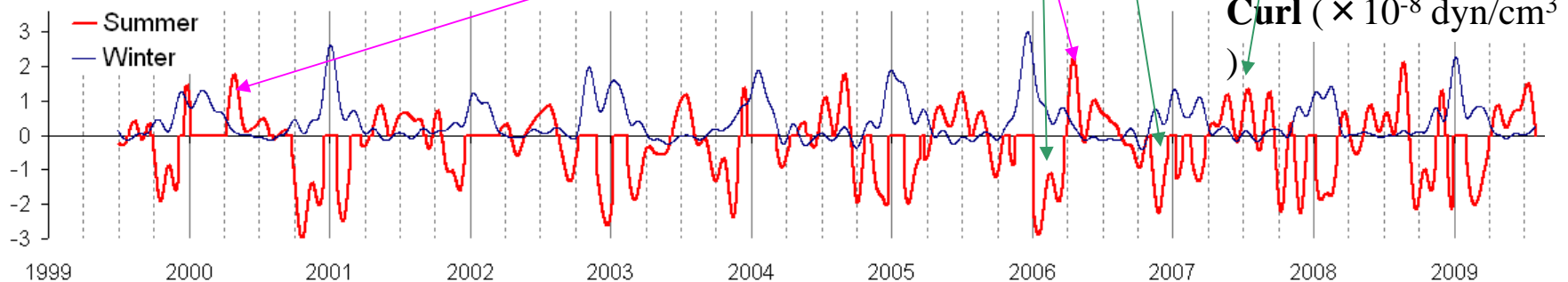


July

Eastern wind,
frequent in late summer

AC curl in late winter and autumn, oscillating between cyclonic and zero (slight AC) curl in late summer.

April 2000, 2006: strong cyclonic curl.

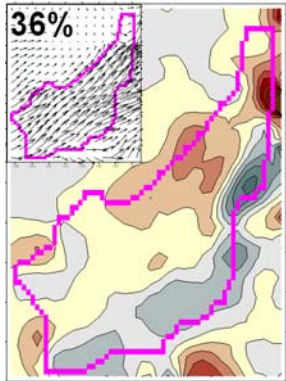


Winter/summer curls for the median spatial EOFs; summer curl when exceeds the winter curl

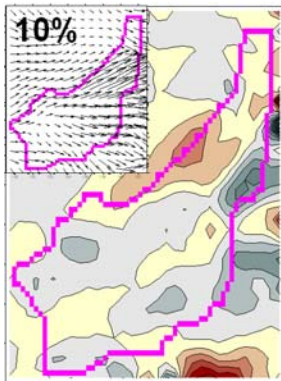
Warm season mode: the northeastern JES

April

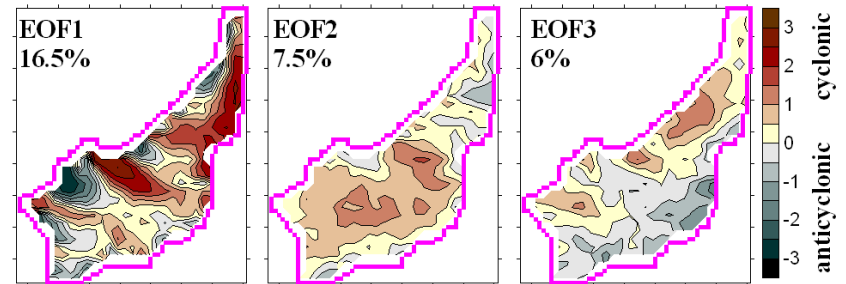
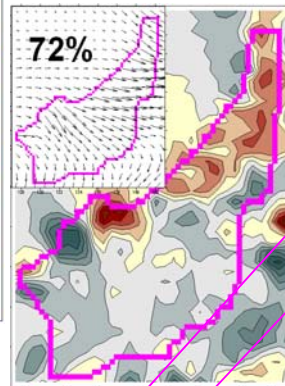
Western or SW
wind, frequent in
spring and autumn



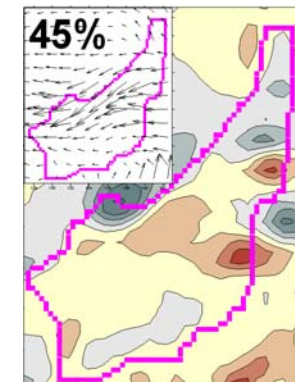
April



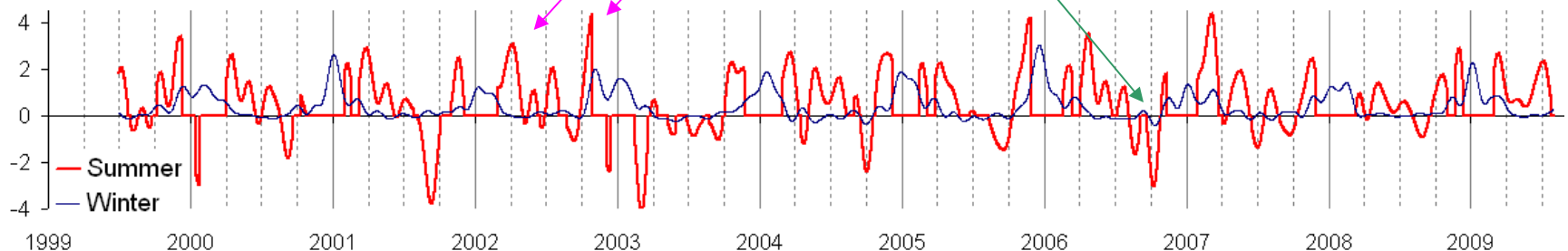
October



Cyclonic curl
in spring/early
summer and
autumn,
frequent AC
curl in late
summer.



July



Winter/summer curls for the median spatial EOFs; summer curl when exceeds the winter curl

Conclusion

- The East Asia Monsoon pattern with characteristic wind directions and seasonal shifts is the leading complex EOF mode of wind stress over the JES.
- Contribution of higher modes is not significant.
- The results are consistent with the previous findings from reanalysis data (Trusenкова et al., 2009).
- Winter pattern is the leading EOF mode of wind stress curl. Fine features are resolved, such as several curl dipoles related to orographic gaps.
- Over the central JES, the AC wind stress curl prevails in late winter and autumn, while oscillations between the cyclonic and weak AC curl occur in summer.
- Over the northeastern JES cyclonic curl prevails in spring and autumn.

Thank you!

