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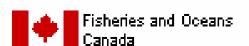
Acknowledgements

Other participants in the **ECOHAB PNW Project**



Related Talks

9:10 MacFadyen 11:30 Peña



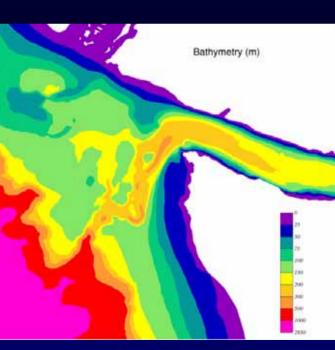
AND ATMOSPH

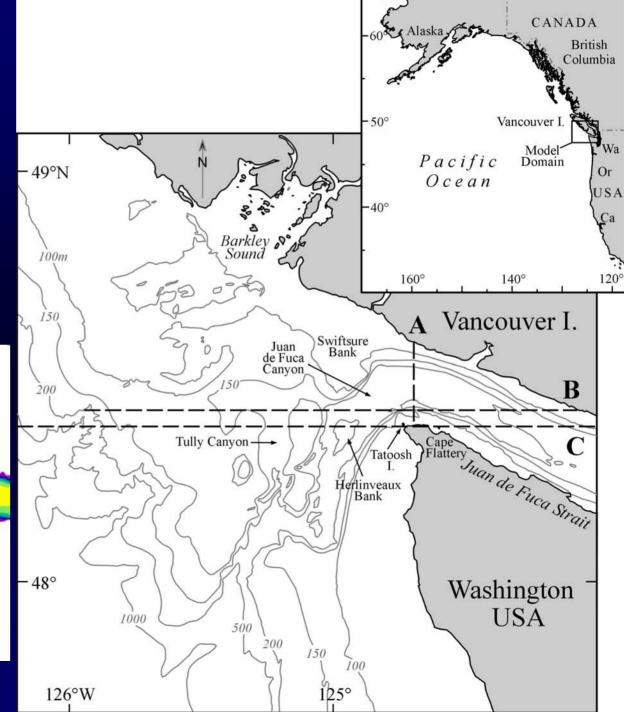
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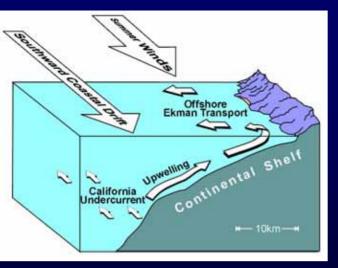
Pêches et Océans Canada



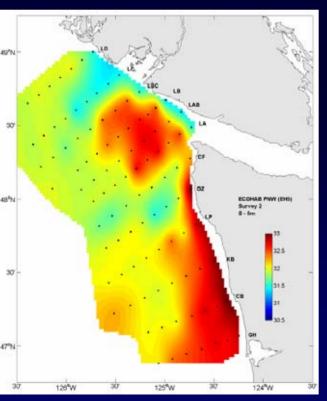
The Region of Interest







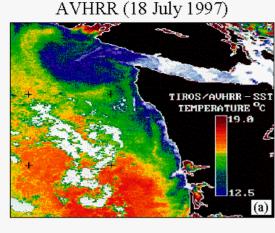
Courtesy of Rick Thomson



Sept 2005 salinity at 5m depth

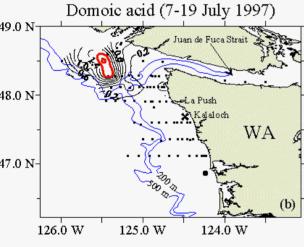
Juan de Fuca (Tully) Eddy

- **summer** upwelling feature off the entrance to Juan de Fuca Strait
- Not classical upwelling, as off WA, OR, CA
- comprised of nutrient-rich California Undercurrent water (Freeland & Denman, 1982) that moves up the Juan de Fuca and Tully Canyons onto the shelf
- Makes the SW Vancouver Island shelf one of most productive fishing regions in the NE Pacific (Ware & Thomson, 2005)









• Eddy seems to be initiation site for toxic Pseudo-nitzschia that can impact clam & crab fisheries along the Washington coast (Trainer et al., 2002)

• ECOHAB PNW:

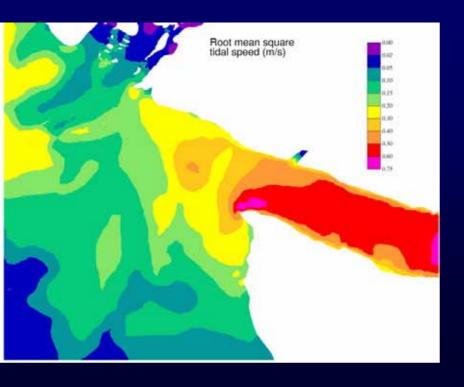
- > multi-disciplinary project to study ecology & oceanography of these harmful algal blooms (HABs)
- Eddy physical dynamics are important for understanding biochemistry and transport of these HABs

- Pseudo-nitzschia cells (7-19 July 1997)
- 48.0 N-47.0 N-

125.0 W

124.0 W

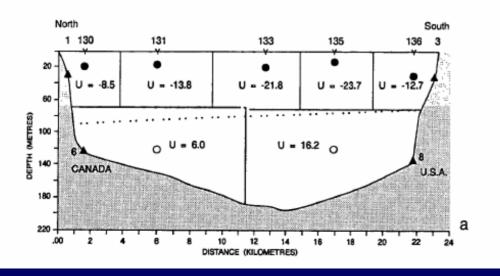
126.0 W

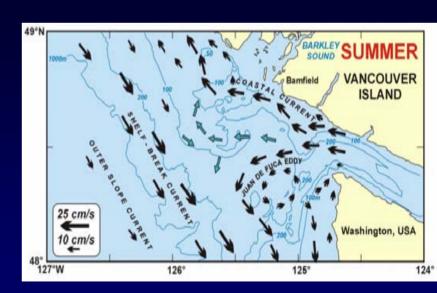


Regional Oceanography

- > Strong tidal, estuarine, & winddriven flows in Juan de Fuca Strait
- > Estuarine flow primarily from Fraser River
- > Summer upwelling winds

380 / A.J. Mark Labrecque, R.E. Thomson, M.W. Stacey and J.R. Buckley

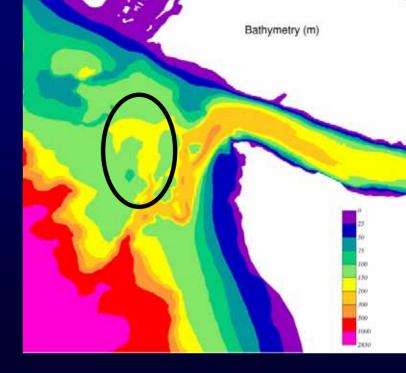


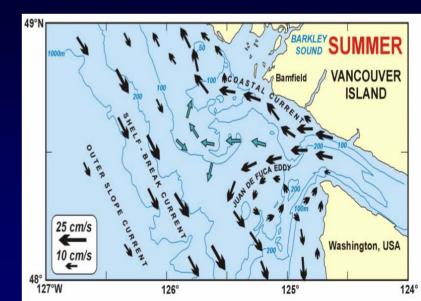


Freeland & Denman (1982) Hypothesis

- eddy closely associated with northern end of Tully Canyon
- > cyclonic eddy spins up when shelf edge currents reverse in spring
- > near surface, inward pressure gradient largely balanced by outward Coriolis force
- > further down water column, Tully Canyon suppresses transverse motions leaving only inward pressure gradient
- pressure gradient forces up-canyon flow, upwelling California Undercurrent water from shelf edge

We can test this hypothesis with numerical model experiments.





Modeling

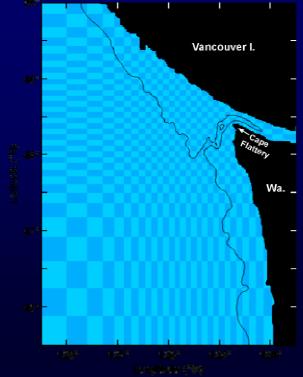
ROMS = Regional Ocean Modeling System

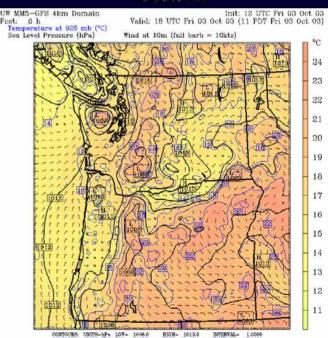
Objectives:

- What combination of winds, estuarine flow & tides are necessary for eddy generation?
- **▶** What are the specific dynamics?
- ➤ How do wind variations affect retention & transport?

Model details:

- > Stretched grid: 1 to 5 km
- ➤ Temperature & salinity initial conditions from summer climatology
- ➤ Average summer winds from UW MM5 atmospheric model (next slide)
- \triangleright M₂, S₂, K₁, O₁ tidal forcing
- > Strong TS nudging at JdF boundary to maintain estuarine flows
- Radiation &/or nudging conditions on N, S, W boundaries
- > No Columbia discharge



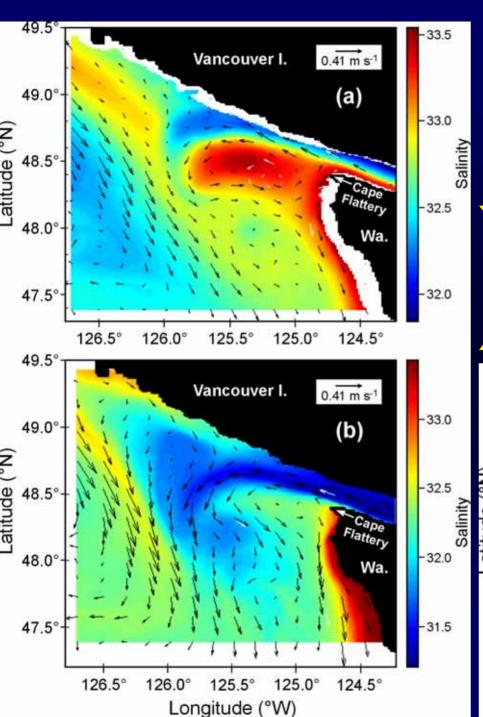


Average Summer Upwelling Winds (m/s). Interpolated from June-Sept 2003-05 MM5 Data. 8 49 48 6 47 Latitude 5 4 45 3 44 2 -128 -127 -126 -125 -124 -123 Longitude

Tinis et al. (2006) compared MM5 winds with buoy data

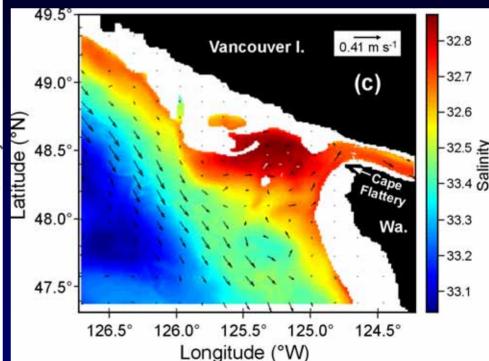
Model Experiments to Study Generation

Experiment	Objective	Initial Conditions	Tides	Estuarine Flow	Winds	Duration
A	Baseline run	Summer climatology	yes	yes	yes	60 days
В	Role of winds	Summer climatology	no	yes	yes	60 days
С	Role of tides	Summer climatology	yes	yes	no	60 days
D	Role of estuarine flow	T and S profiles	yes	no	yes	60 days



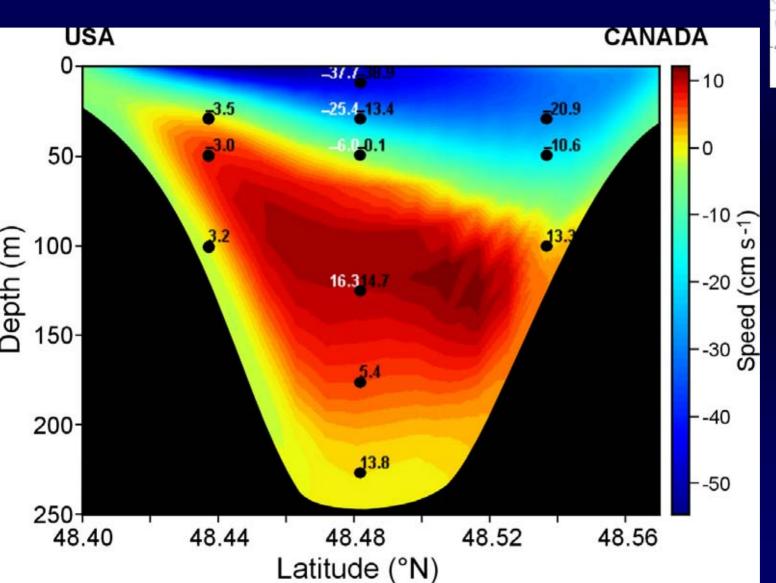
Baseline Run Validation:

Average (days 46-60) flows & salinity at 0, 35, 100 m depths

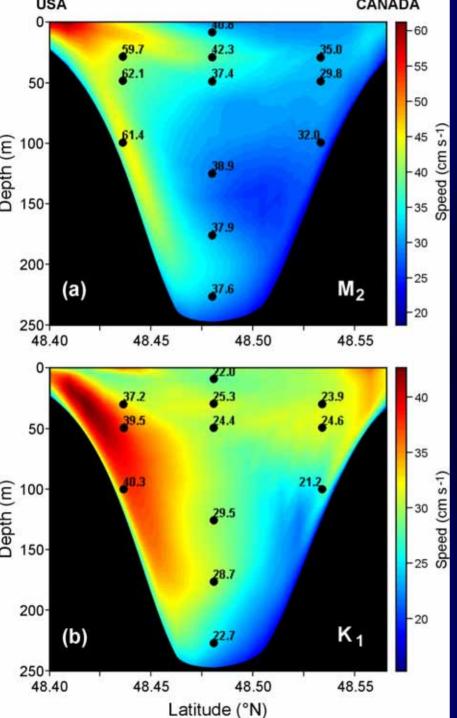


Baseline Run Validation:

Mean flows in Juan de Fuca Strait

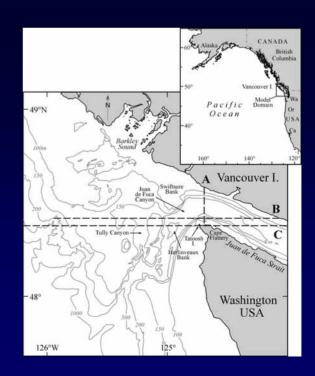


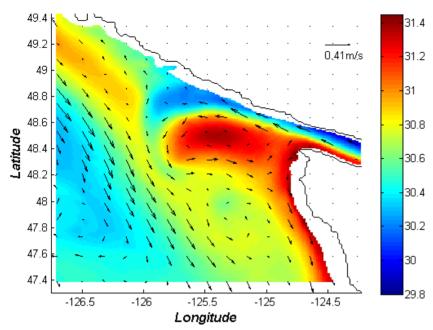




Baseline Run Validation:

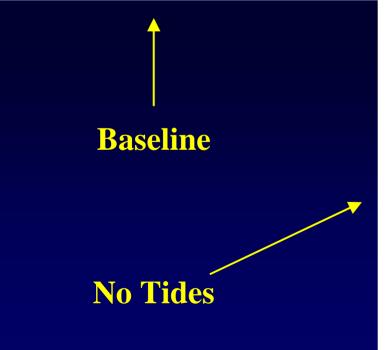
 $\begin{array}{c} M_2 \ \& \ K_1 \\ \text{along-strait maximum speeds} \\ \text{(major semi-axes cm/s)} \end{array}$

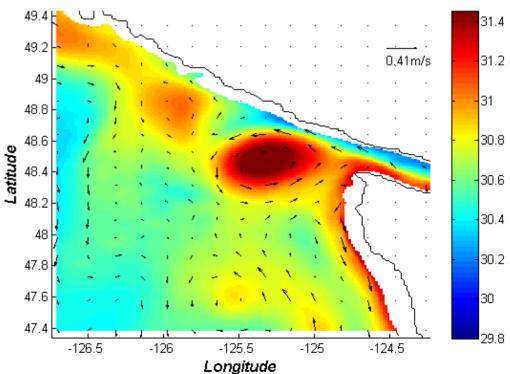


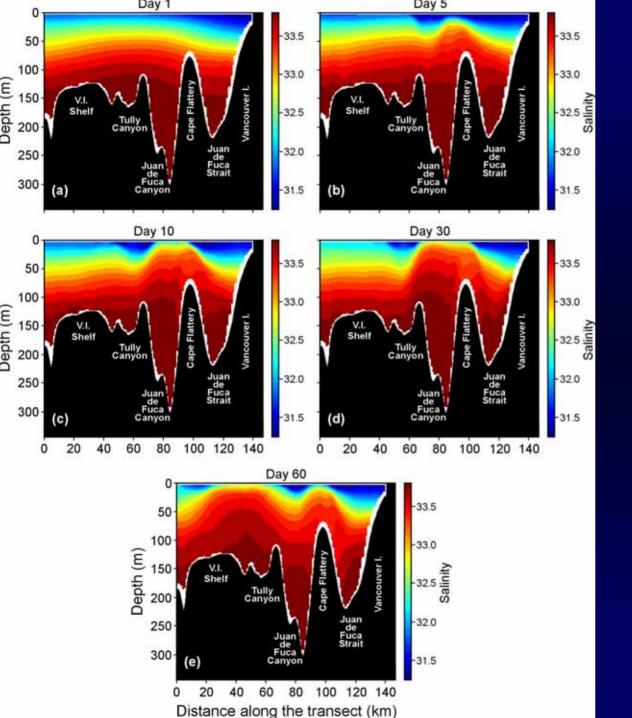


No Tides Run

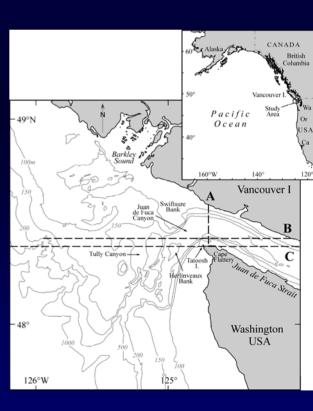
Average (days 46-60) flows & salinity at 35 m depth



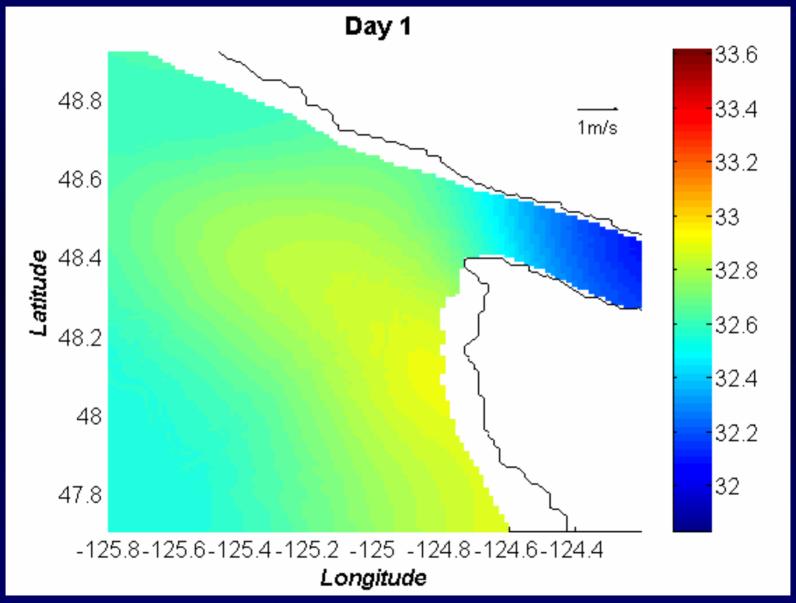




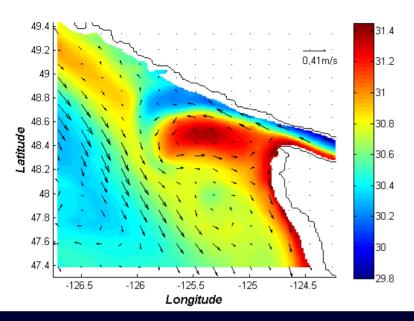
Eddy Development with No Tides



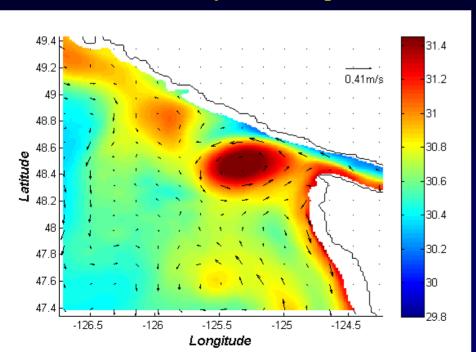
Eddy Development with No Tides



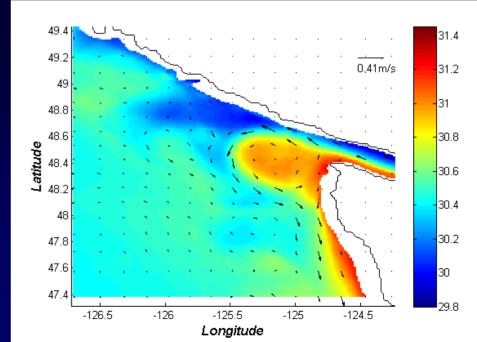
Daily 35m salinity and velocity

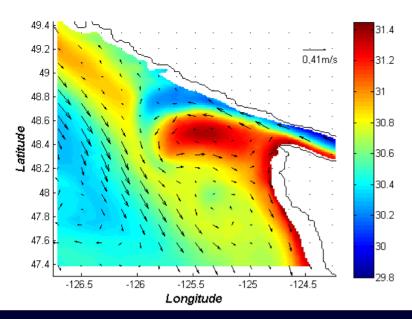


Average (days 46-60) flows & salinity at 35 m depth

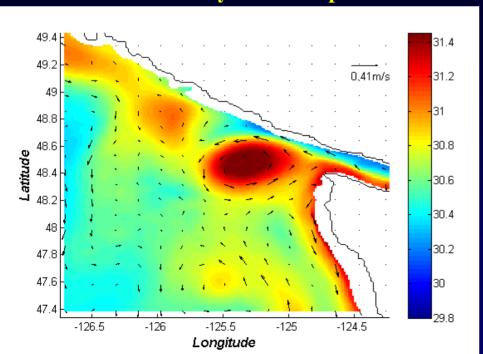


Baseline vs No Tides vs No Winds

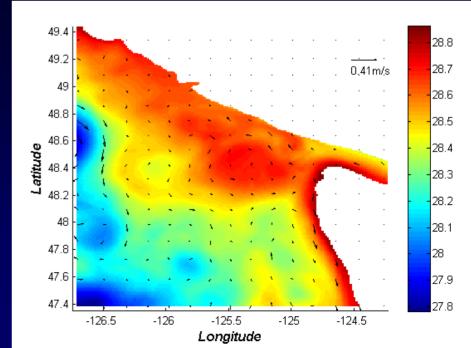




Average (days 46-60) flows & salinity at 35 m depth

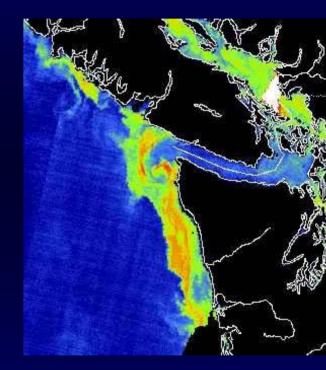


--- Baseline vs No Tides vs No Estuarine Flow



Summary of Numerical Experiments

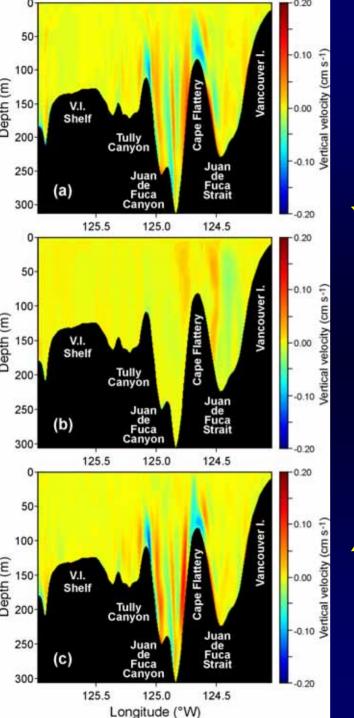
- Eddy formed from enhanced upwelling off Cape Flattery
- Winds, tides & no estuarine flow
 - A very weak eddy
 - > Comparable to upwelling along coasts
- Winds &/or tides with estuarine flow
 - Stronger eddy
 - Wind-driven eddy stronger than tidal-driven



MERIS chlorophyll image: June 3, 2003 Courtesy of Jim Gower & Steph King

What causes the:

- i. Enhanced upwelling off Cape Flattery?
- ii. Cyclonic eddy formation?



Upwelling Mechanisms:

1. Mean vertical velocity (along transect B)

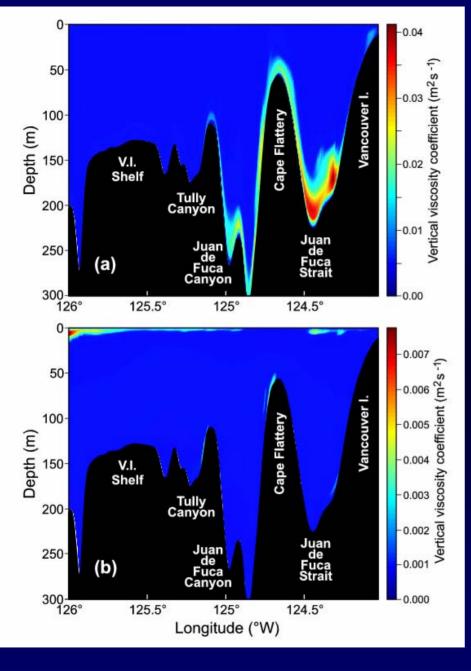
Baseline

No Tides

No Winds

➤a) & c) suggest vertical tidal residuals: Garrett & Loucks (1976)?

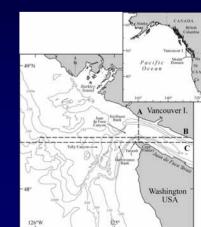


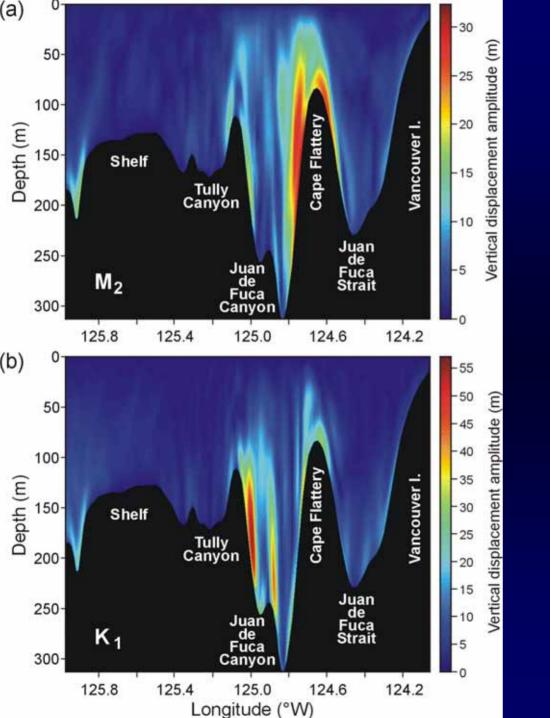


Why stronger eddy with no tides?

Winds & tides

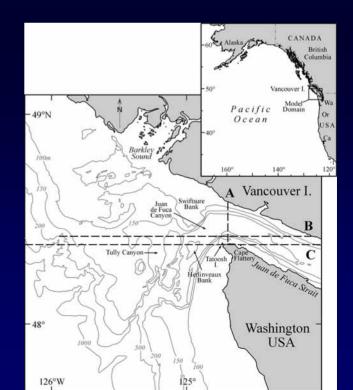
Winds but no tides

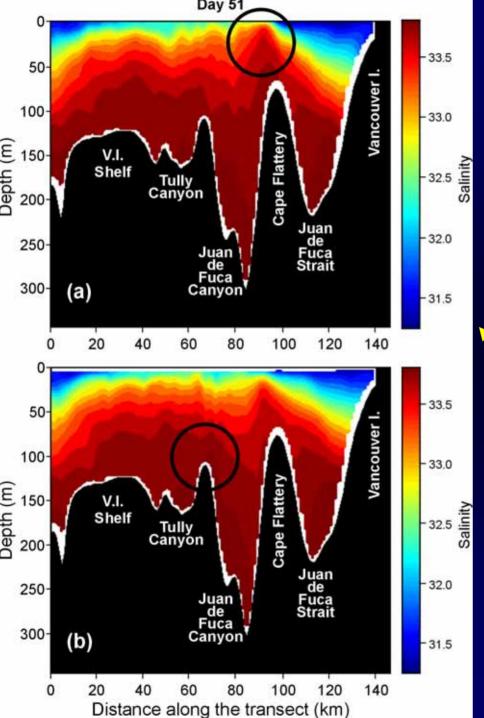




Upwelling Mechanisms:

2. Tidal vertical displacement amplitudes (along transect B)





Upwelling Mechanisms:

3. Cross canyon waves (internal tides?)

Salinities across transect B

hour 23 = flood tide



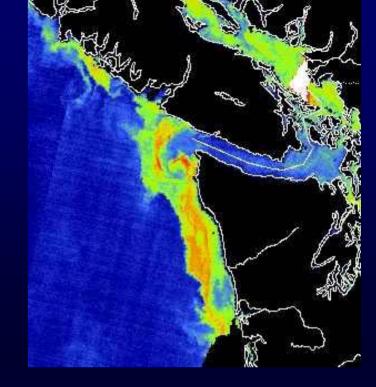
Why

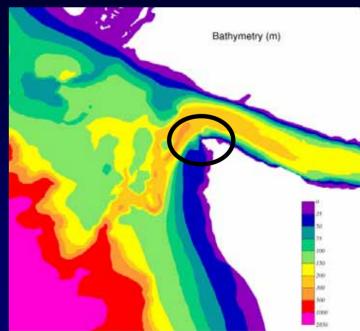
i) a cyclonic eddy?

- upwelled denser water causes a surface depression & inward pressure gradients
- ➤ geostrophy → cyclonic eddy

ii) enhanced upwelling off Cape Flattery?

- proximity of dense water in bottom estuarine flow
- > 200m depth contour only 4km away





Relevance to HAB Development & Transport

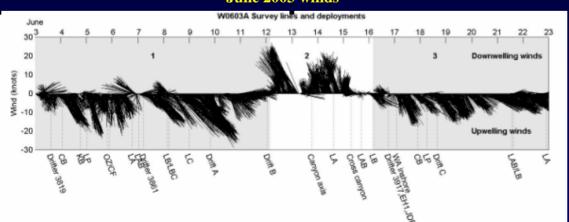
- Eddy can retain *Pseudo-nitzschia* & provide nutrients for them to grow "bio-reactor"
 - low iron/copper availability seems to correlate with high toxin concentration
- Variations in winds (& tides) will affect phytoplankton retention & transport
 - MacFadyen et al.

June 6

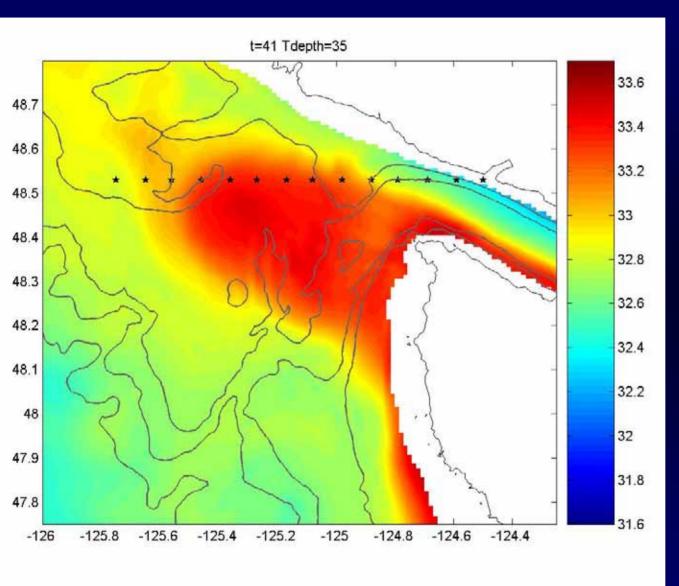
June 15

MERIS images courtesy of Jim Gower & Steph King

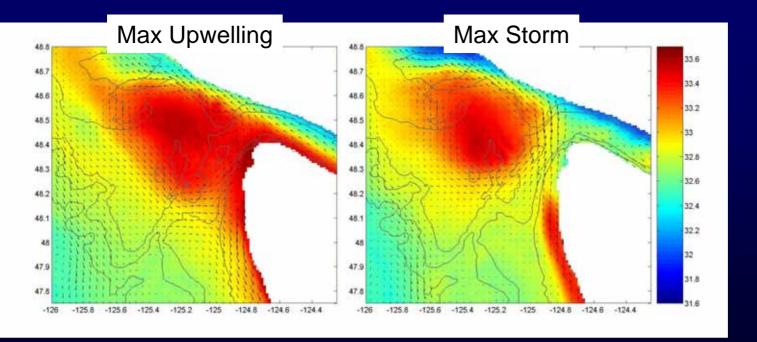
June 2003 winds



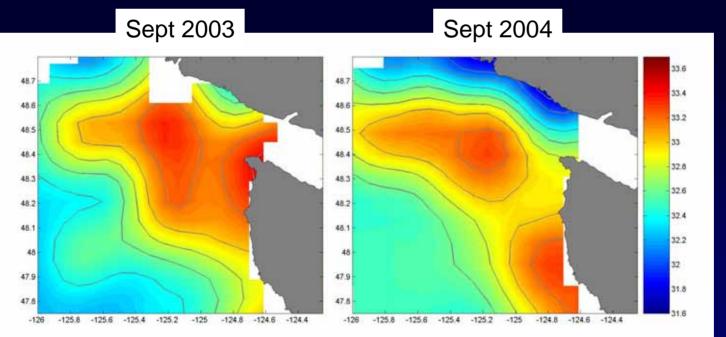
ROMS Variable Wind Forcing Experiments



- initialized from 40day run with average upwelling winds
- wind forcing either strong upwelling or storm, ramped up over 5 days, returned to average over next 5 days
- near-surface drifters "deployed" in locations shown



ROMS

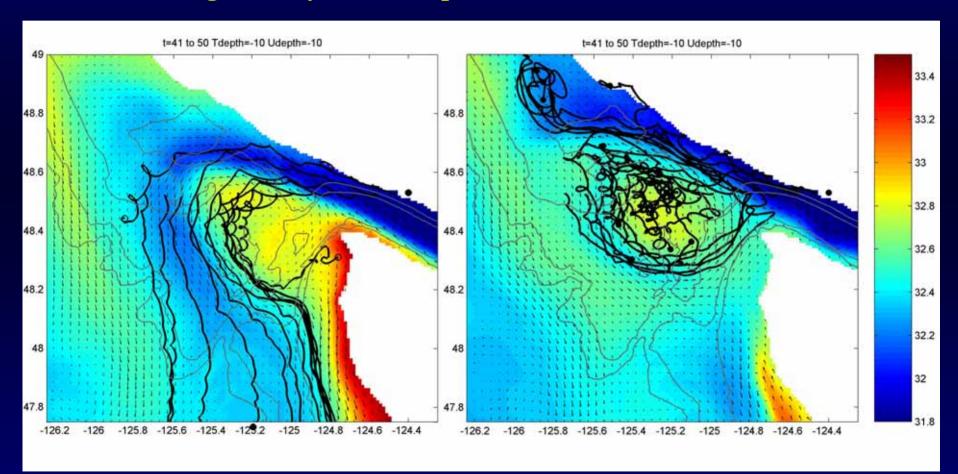


ECOHAB PNW surveys

Transport under Wind-driven Variability

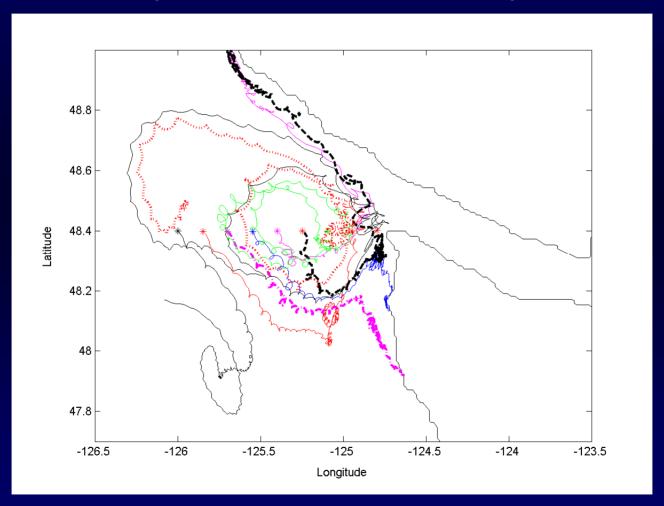
ROMS run with time varying winds – initialized from 40-d previous run

Average salinity over 10-d period with surface drifter tracks



Transport under Wind-driven Variability

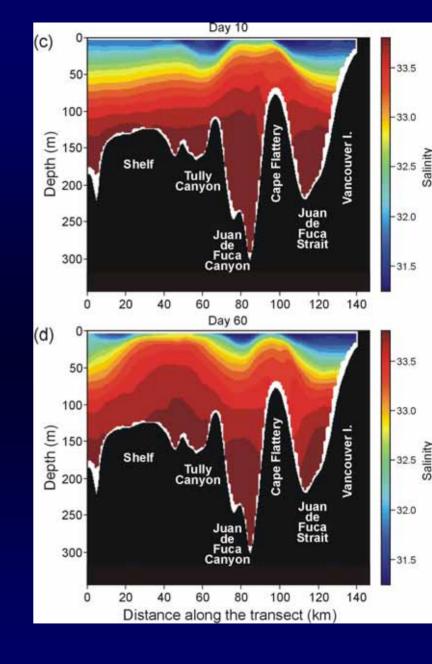
Mean summer wind (days 1-30), 50 % stronger summer wind (days 31-35), winter wind (days 36-40)



9 particles released at 35m depth at day 5

Summary & Conclusions

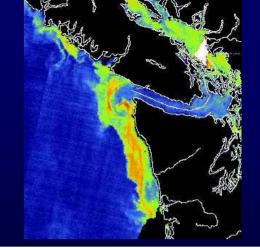
- Model runs suggest Juan de Fuca Eddy is generated by enhanced upwelling off Cape Flattery (Allen, 2000)
- Not deep flows up Tully Canyon & onto shelf (Freeland & Denman, 1982)
- *However*, Tully Canyon flows may help sustain eddy after it forms & moves westward

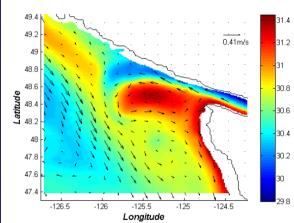


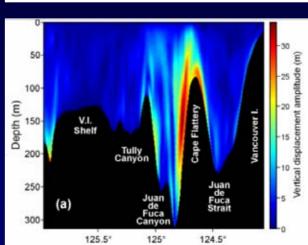
No tides simulation

Summary (cont'd)

- Reasonable agreement between summer observations & model
 - > Confidence in model dynamics
- Eddy generated with estuarine flow and upwelling winds &/or tides
 - Key: proximity of dense water to Cape Flattery
 - > Bottom estuarine flow in strait & canyon
- Tidal upwelling arises in three ways
 - 1. Vertical displacements off Cape Flattery (during flood tides)
 - 2. Vertical tidal residual flows
 - 3. Sloshing across canyon bottom & onto Vancouver Island shelf during ebb tide
 - > Internal tides?







Summary (cont'd)

- Variations in wind will affect eddy retention and transport
 - > Storms after average upwelling strengthen retention & push particles northward
 - Strong upwelling pushes particles to south
 - Strong upwelling followed by storm can push particles to Washington coast & impact shellfish
- Climate change & the eddy (FUTURE)
 - > Still have tides & wind
 - Estuarine flow may change
 - **▶** Upwelled water properties (CUC) may change
 - **Less productive ?**
 - More work needed

